



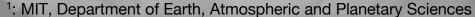


How do rocks fail? Strain localization and microstructural transformations in the earthquake cycle

Matěj Peč

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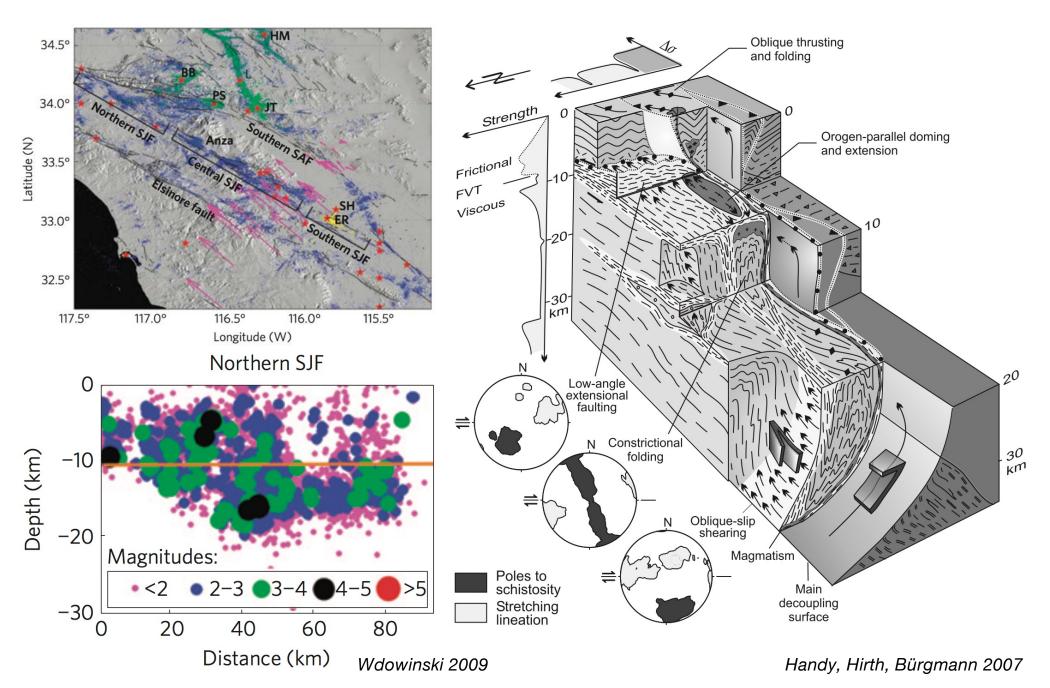




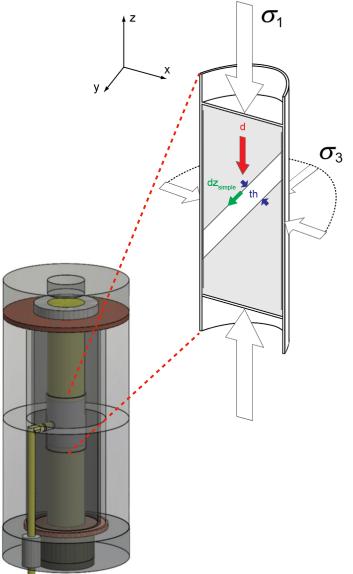
^{3:} Utrecht University, Department of Earth and Planetary Sciences

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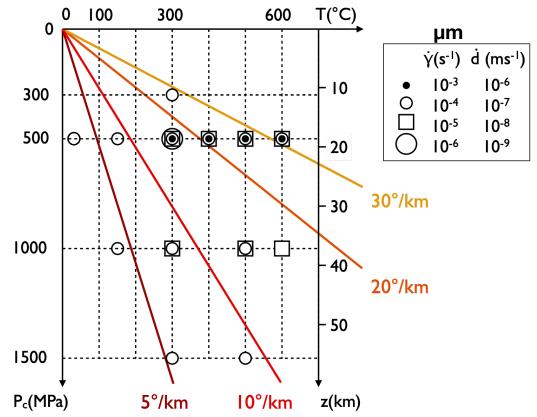
Fault Zones



Experiments on cataclasites



- Verzasca gneiss granitoid
 - ~ 37% Qtz, 33% Plg, 28% Kfs, 2% micas
 - Initial grain size, $d \le 200 \mu m$, 0.2 wt% H₂O added
 - Newest additions: **AE** monitoring, **local T** measurements

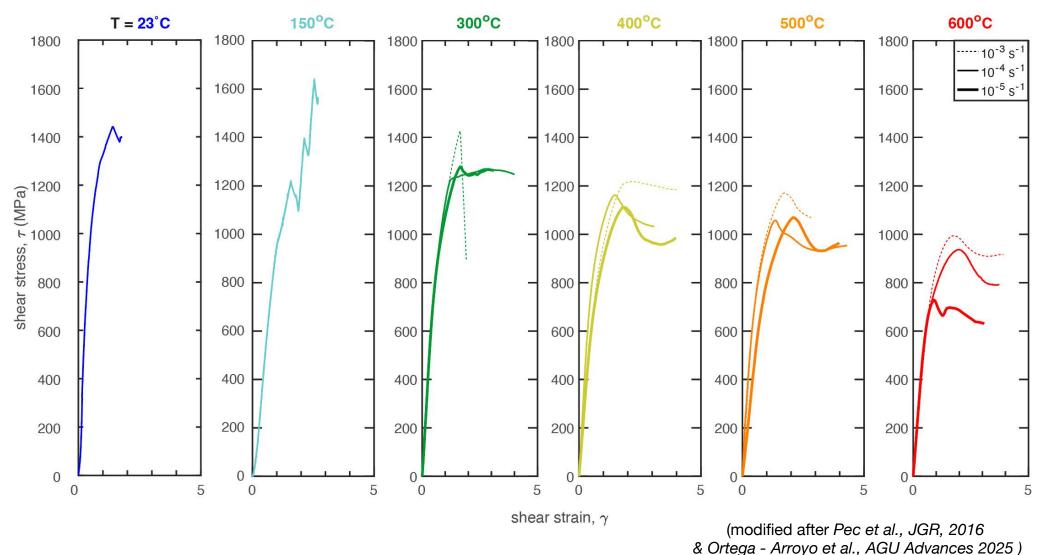


Modified after Précigout et al. 2018

(modified after Pec et al., JGR, 2016)

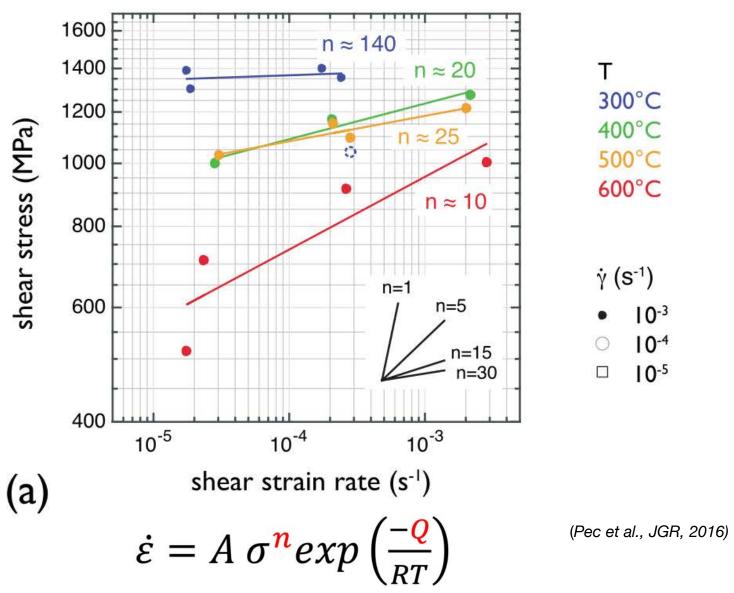
Mechanical Data

Pc = 500 MPa



- Transition from stick-slip to slow slip with increasing T and decreasing rate
- Weak temperature & rate dependence

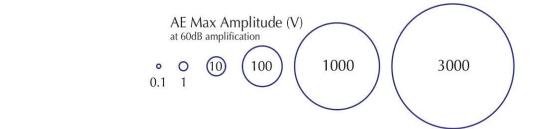
Mechanical Data

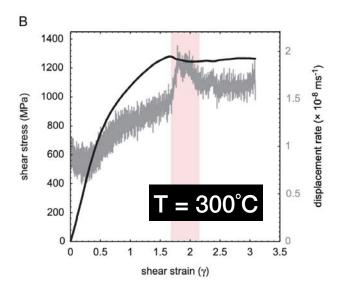


Stress exponent decreases with increasing T

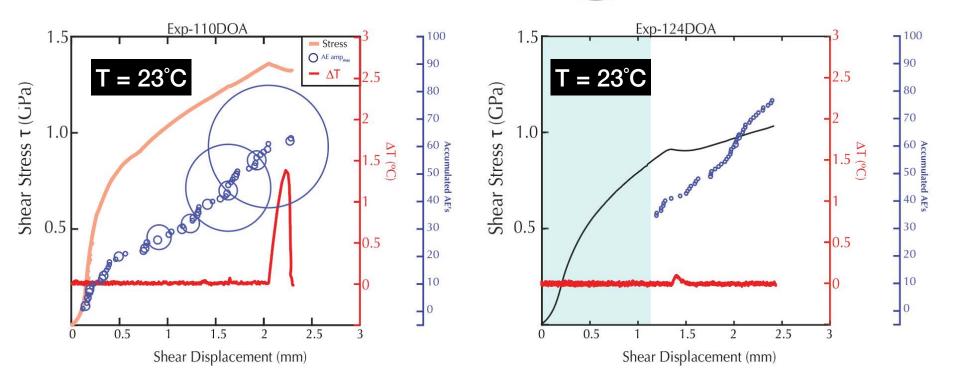
Labquakes & slow slip

- Abrupt failure is accompanied by rapid acceleration, T spikes & audible AE
- Slow weakening is accompanied by modest acceleration, smaller and longer duration T transients





Pec et al. 2012, EPSL

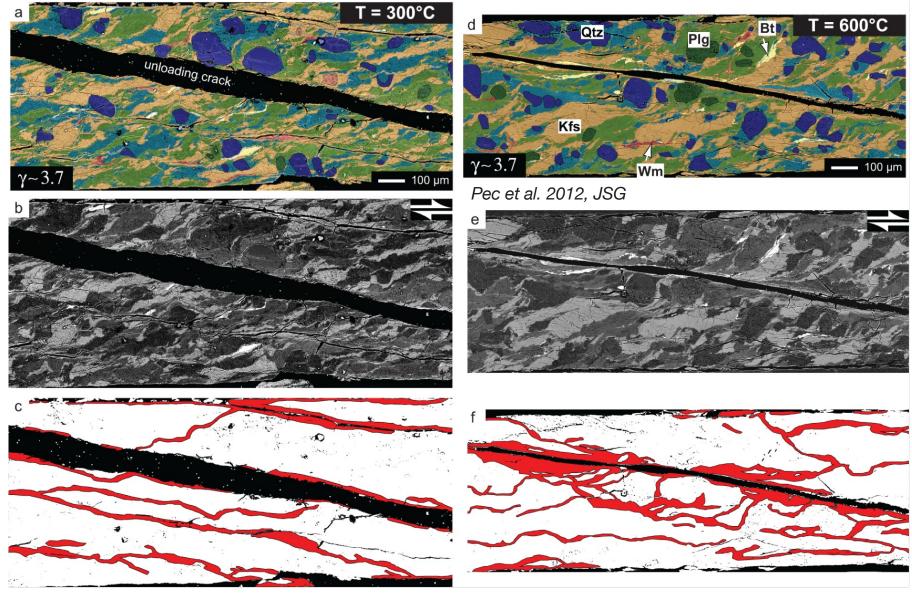


Ortega-Arroyo et al., 2025, AGU Advances & in prep.

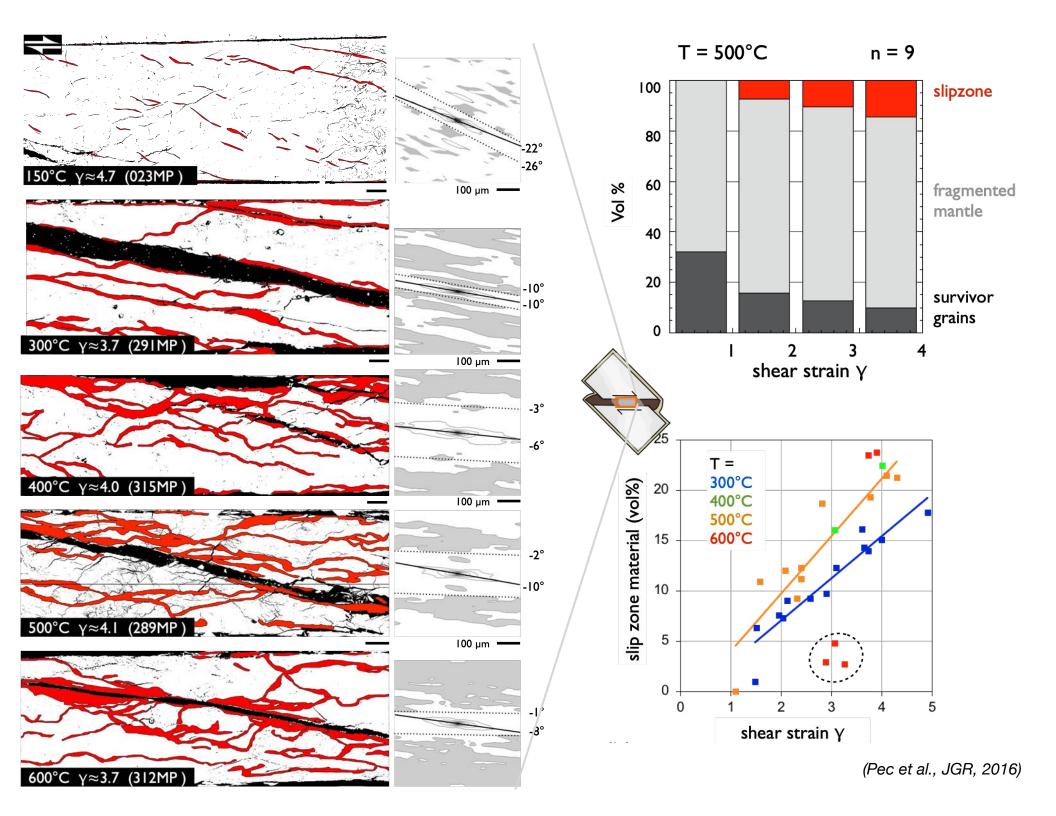
Microstructures

Localization into S-C-C' structures





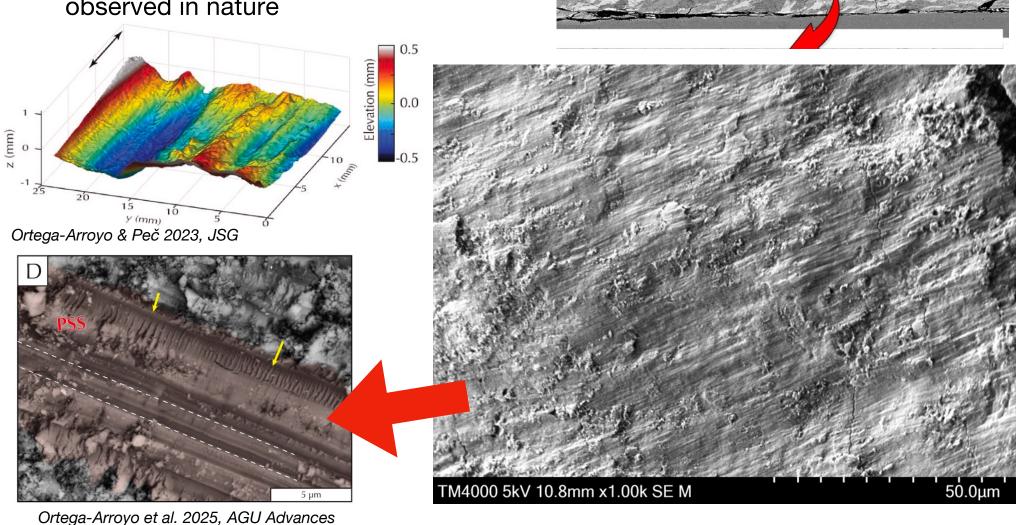
Pc ~500 MPa, γ ~ 3.7, γ ~10⁻⁴ s⁻¹



Microstructures

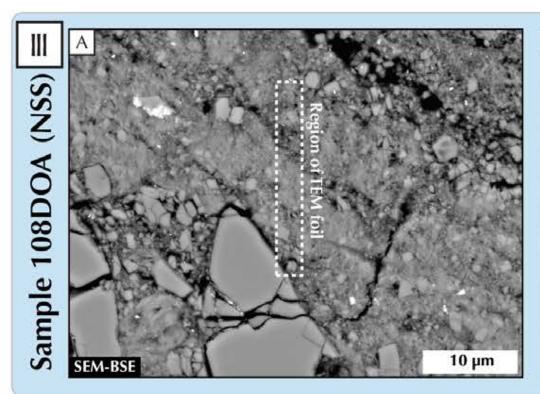
 Unloading cracks exploit the slip zones and expose a "fault surface"

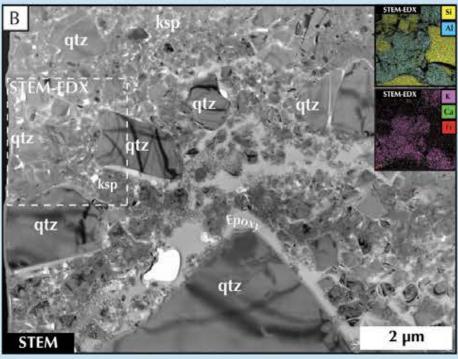
 Qualitatively similar to slickensides observed in nature



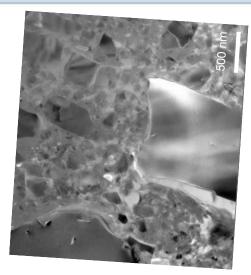
Slip zone origin: Comminution

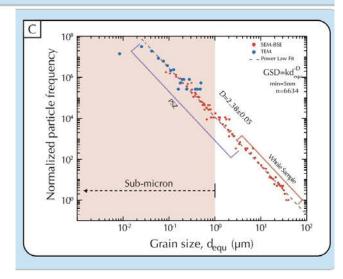
Ortega-Arroyo et al. 2025, AGU Advances



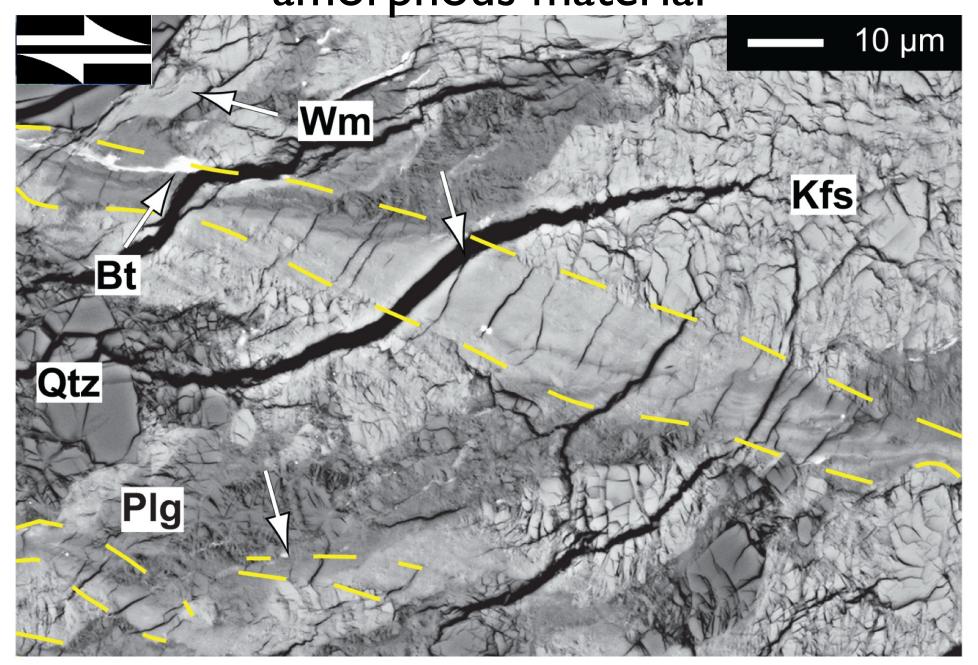


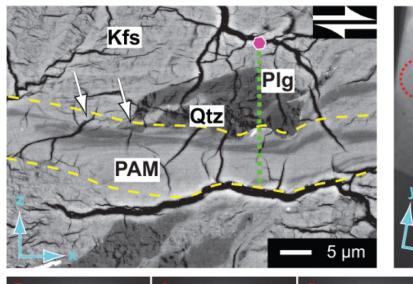
- Wide spread comminution to nm sizes
 - Smallest grains ~ 5 nm
- Qtz strong, Fsp weak
- No evidence of "grinding limit"

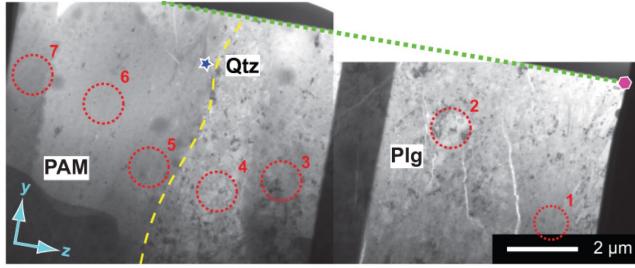


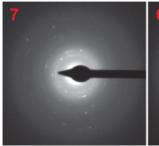


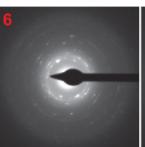
Slip zone origin: Transformation to partly amorphous material

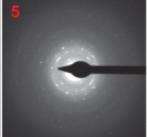


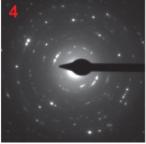


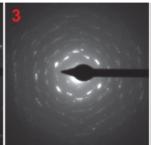


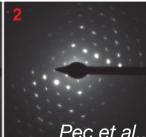


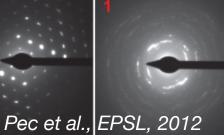


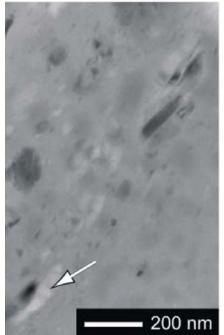




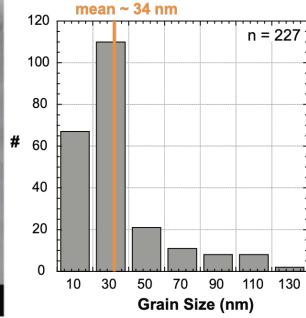






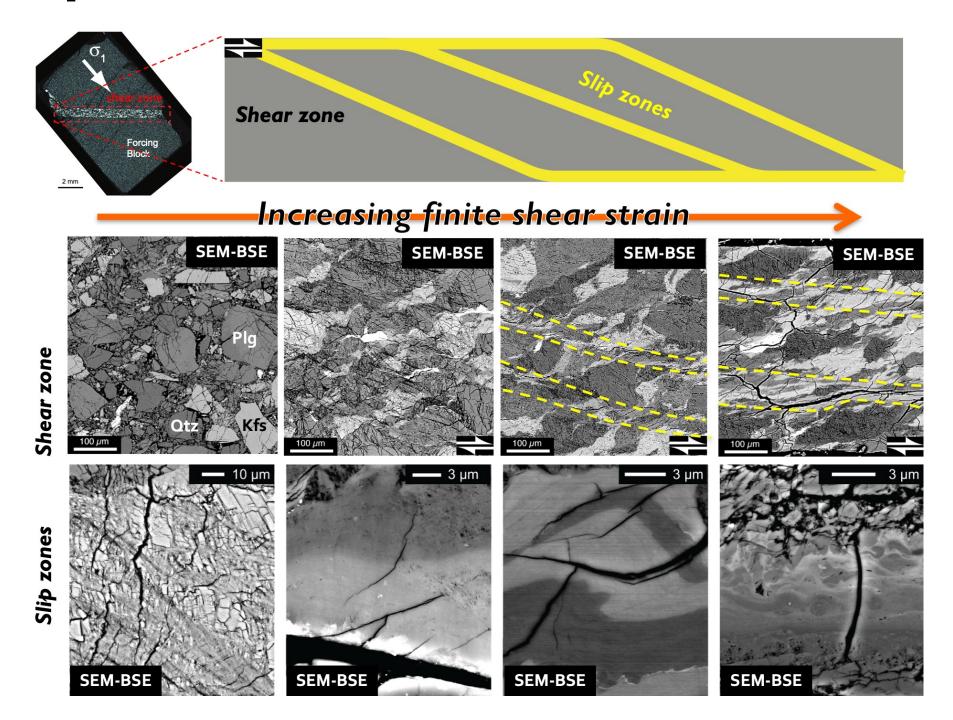






- Up to ~ 90% of PAM is TEM-amorphous
- smallest detectable grains, d ≈ 8 nm (volume ~ 370 unit cells, out of which ~200 unit cells form the surface)
- γ slip zone > γ bulk

Slip zone evolution

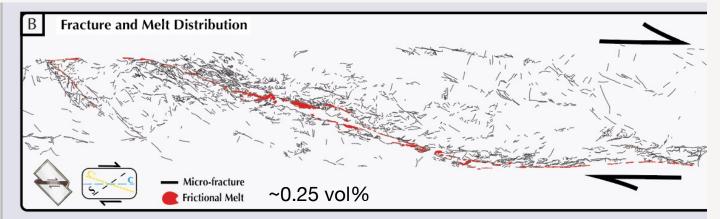


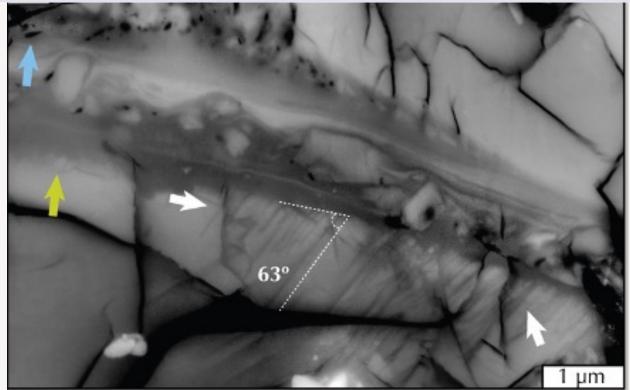
Slip zone evolution: Abrupt Failure

Far-Field at 1Hz -50 1.3 -100 dx_{shear} (mm) 0.2 0 3

 $\Delta \tau(MPa)$

Microstructural Evidence for Heating





Ortega-Arroyo et al. 2025, AGU Advances

Time (s)

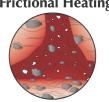
0

-30

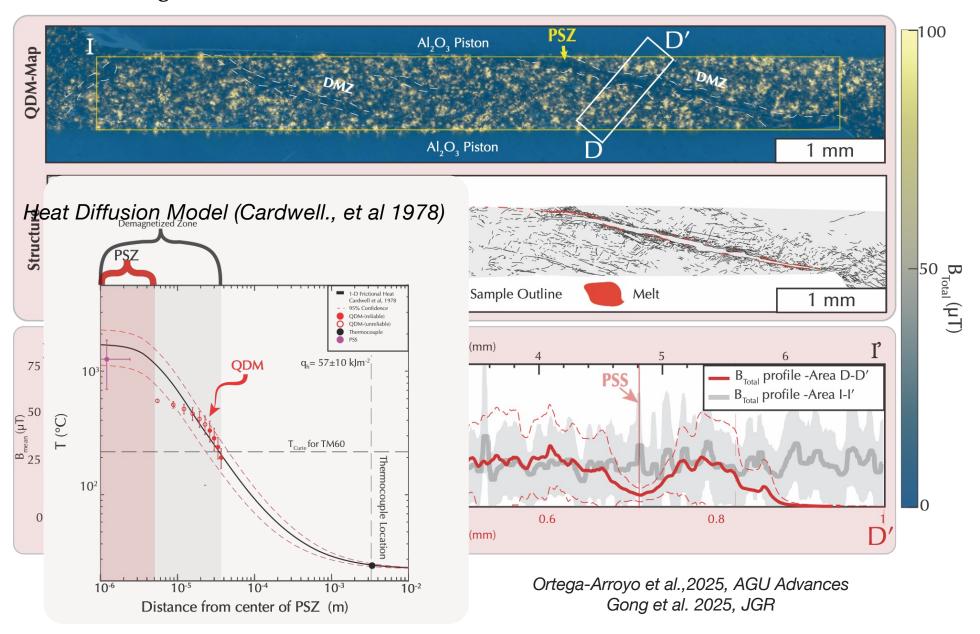
0.9

30

Quantum Diamond Microscopy



Demagnetized Zones (DMZ) around PSZ with melt



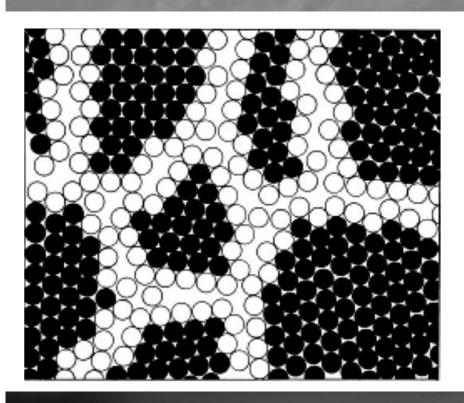
^{2 µm} Slip in fault rocks is localized in thin volumes containing sub-micron materials over a broad range of conditions

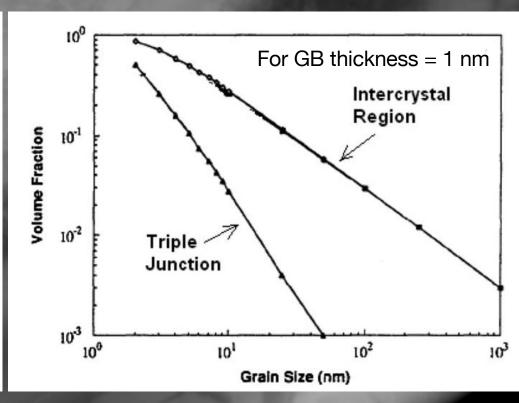
- I) Occur both during **stable sliding** as well as during **stick-slip**
 - 2) Localize strain, i.e. are **weaker** than the surrounding catclazites
 - 3) Stick-slip rupture exploits these zones

1 µm

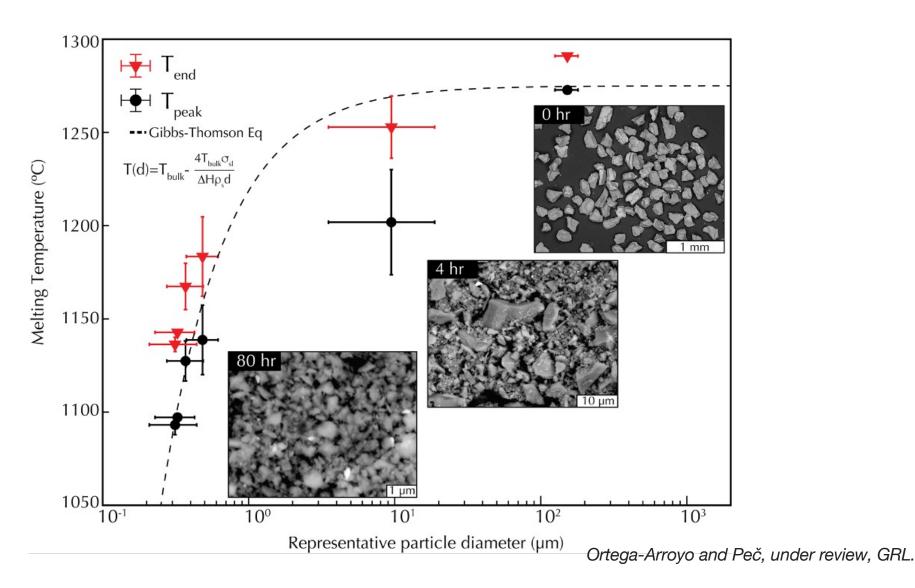
What makes nano-crystalline materials "different"?

- Surface dominated materials
- Physical properties? Rheology?



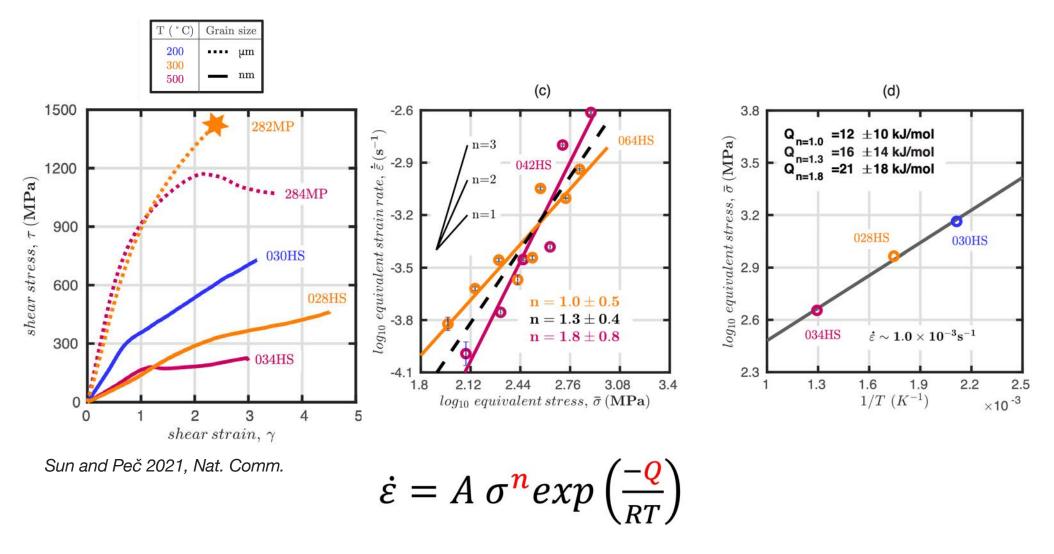


Melting of nano-crystalline fault rocks



- Nanocrystalline materials exhibit a profound melting point depression
- Vulnerable to T-induced weakening

Rheology of nano-crystalline fault rocks

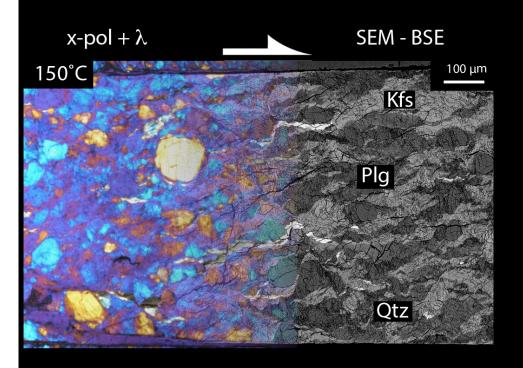


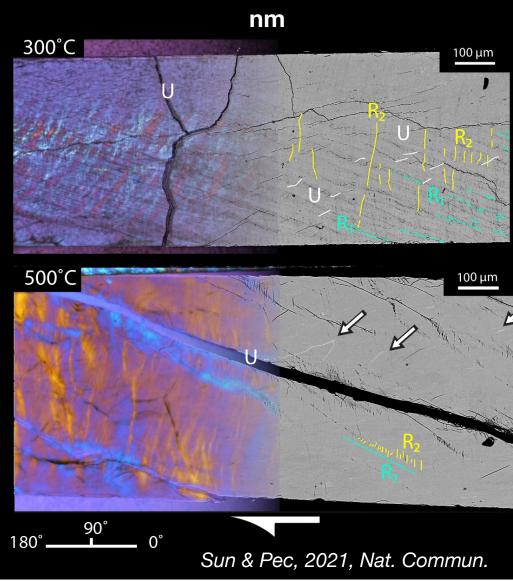
- Nanocrystalline (d \approx 100 nm) materials **produced** by planetary ball milling **in bulk**
- ~ order of magnitude weaker than µm cataclasites under identical conditions
- •Strongly rate strengthening (n = 1-2), low T sensitivity ($Q \approx 10 20 \text{ kJ/mol}$)

Microstructures

- nm-fault rocks are optically and mechanically anisotropic (?!?)
- Sub-wavelength grating effects?

μm

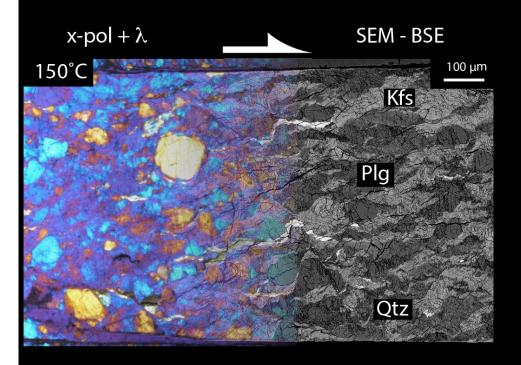


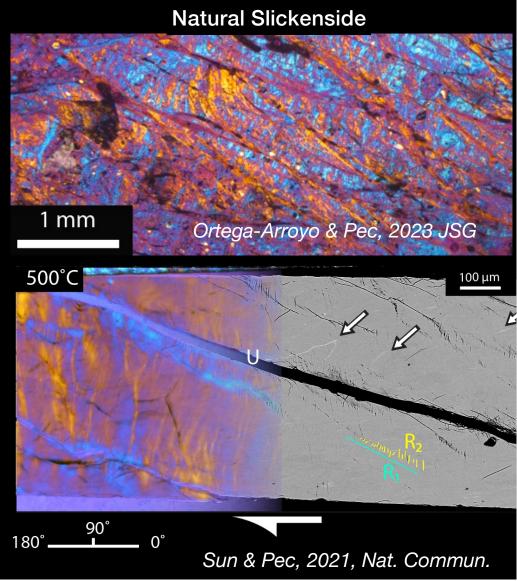


Microstructures

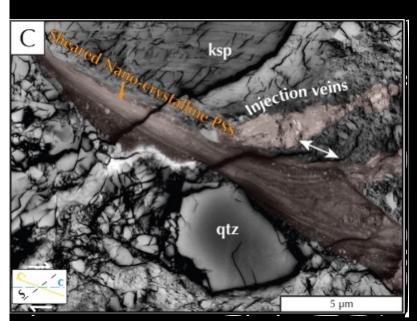
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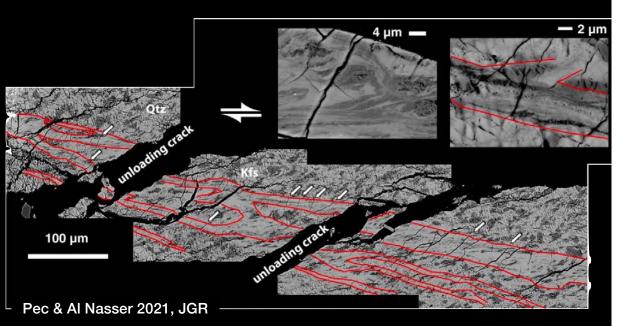
μm





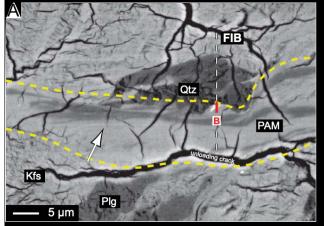
How do Rocks Fail?

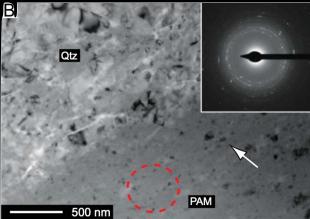


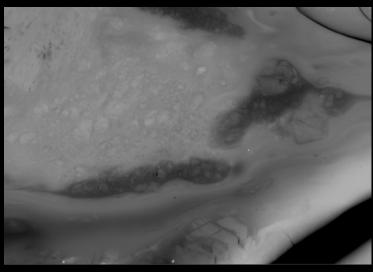


Ortega-Arroyo et al., 2025, AGU Advances

 $Pc = 300 - 1500 MPa, T = 21 - 750^{\circ}C$







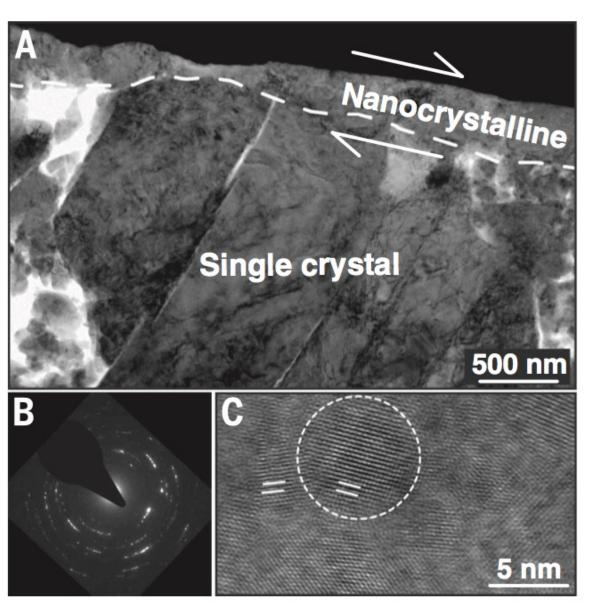
- 1 μm Pec et al. 2016, JGR

Pec et al. 2012, EPSL

Low P-T experiments

 $Pc = 50 MPa, T = 18 - 140^{\circ}C$

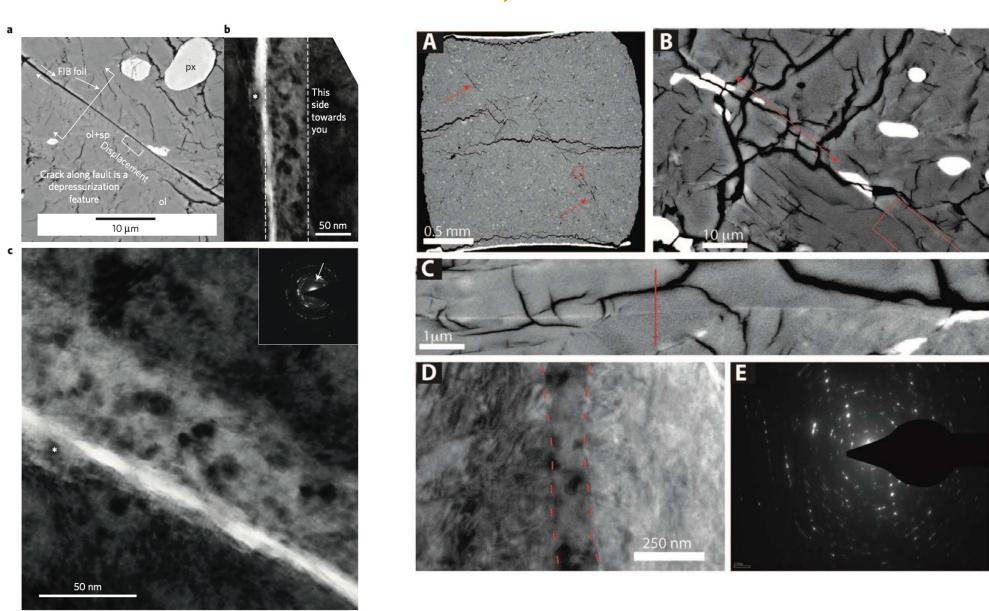




Verberne et al. 2014

Ultra-high P-T experiments

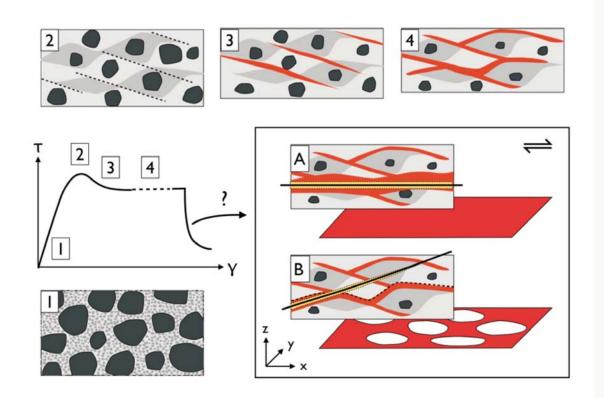
 $Pc = 1.3 - 5 GPa, T \approx 750 - 950^{\circ}C$



Green et al. 2015

Schubnel et al. 2013

How do Rocks Fail?



Pec et al. 2016

Takeaways

- Sub-micron fault rocks are ubiquitous in zones of extreme localization
- nm- surface-dominated materials have distinct properties from µmvolume-dominated materials
- nm-fault rocks are ratestrengthening, with low temperature sensitivity and display a melting point depression of hundreds °C
- While nm-fault rocks are ratestrengthening, they are also intrinsically very weak (≈ order of mag. lower viscosity) and vulnerable to melting
- Percolation of weak material –
 accelerating slip at constant T and stress,
 positive feedback with T increase.
- Analogous to deep-focus EQ?

Thank you!

