

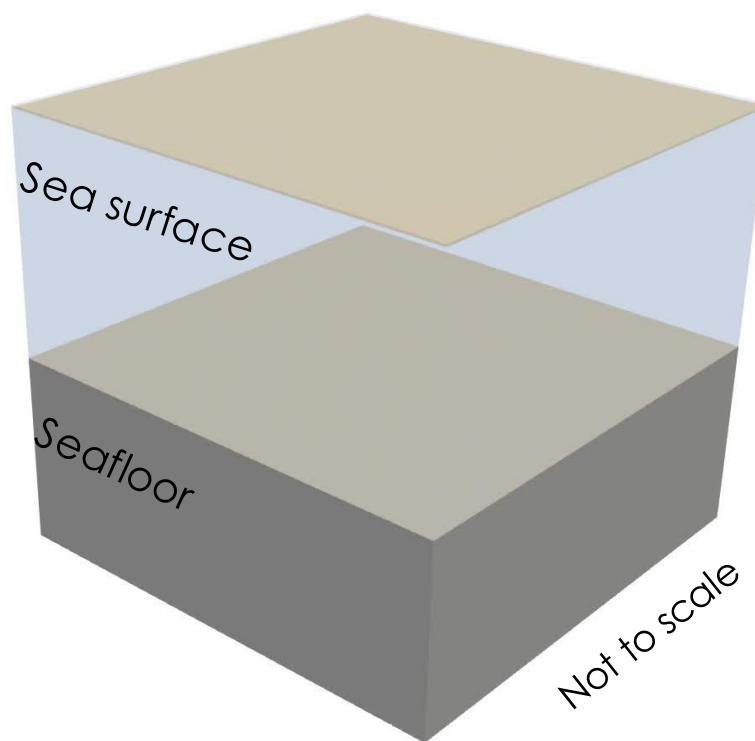


3D fully coupled earthquake dynamic rupture and tsunami benchmarks with varying bathymetric complexity



Time: 0.00 s

Tsunami/Z-disp. (m)



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**2025 SCEC Dynamic
Rupture Workshop**

December 2nd, 2025



UC San Diego



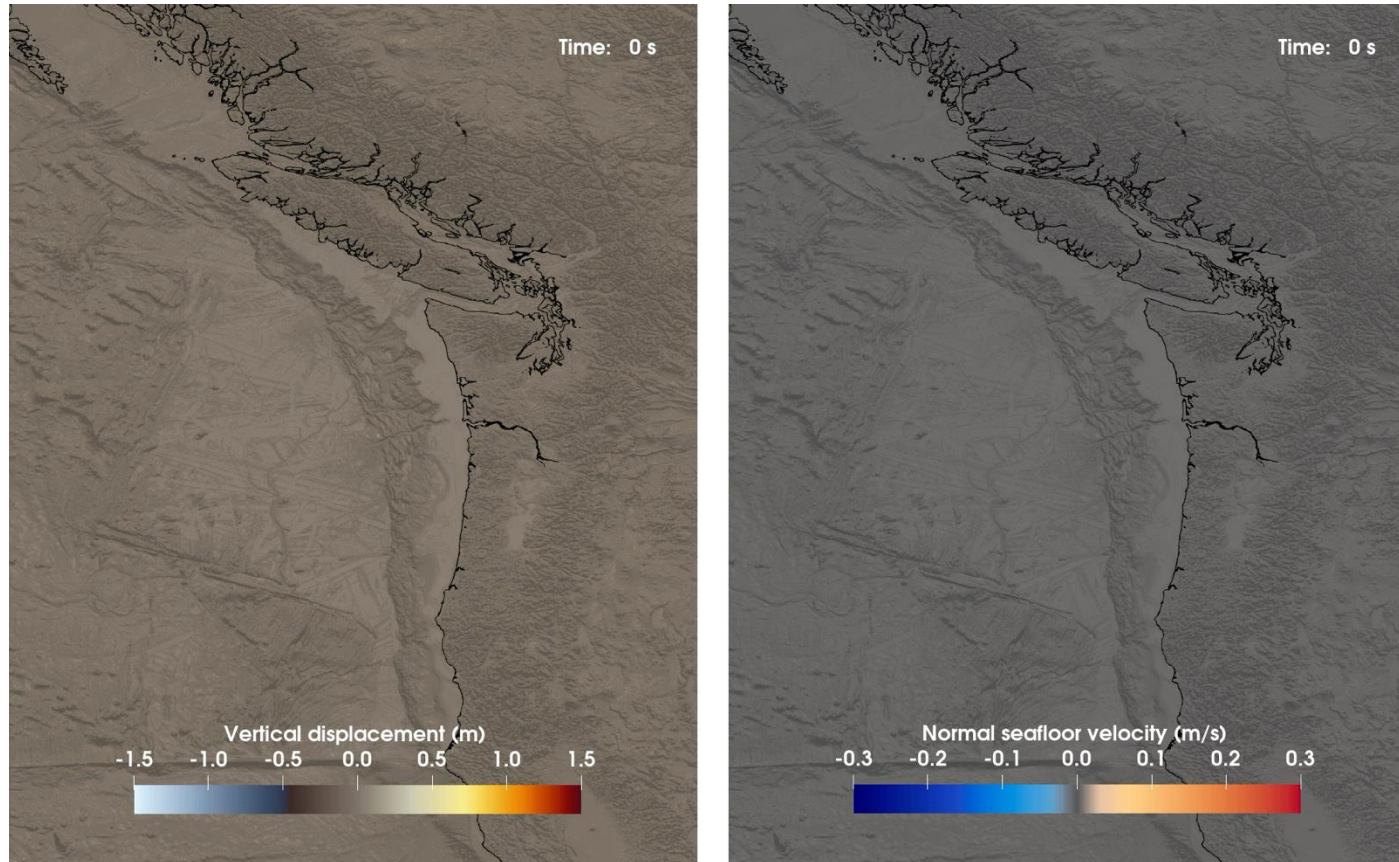
The Tsunami Problem Version 1 (TTPV1)



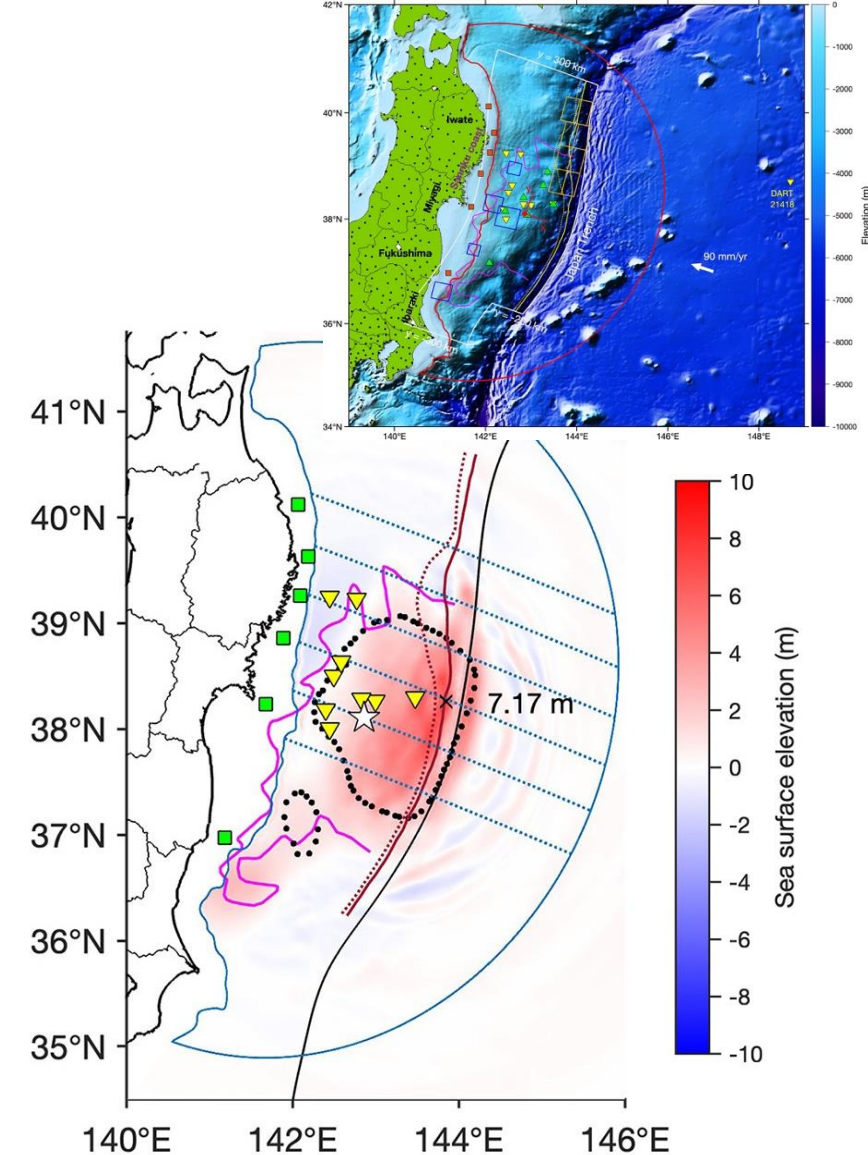
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Realistic earthquake-tsunami scenarios help to constrain the tsunami hazard

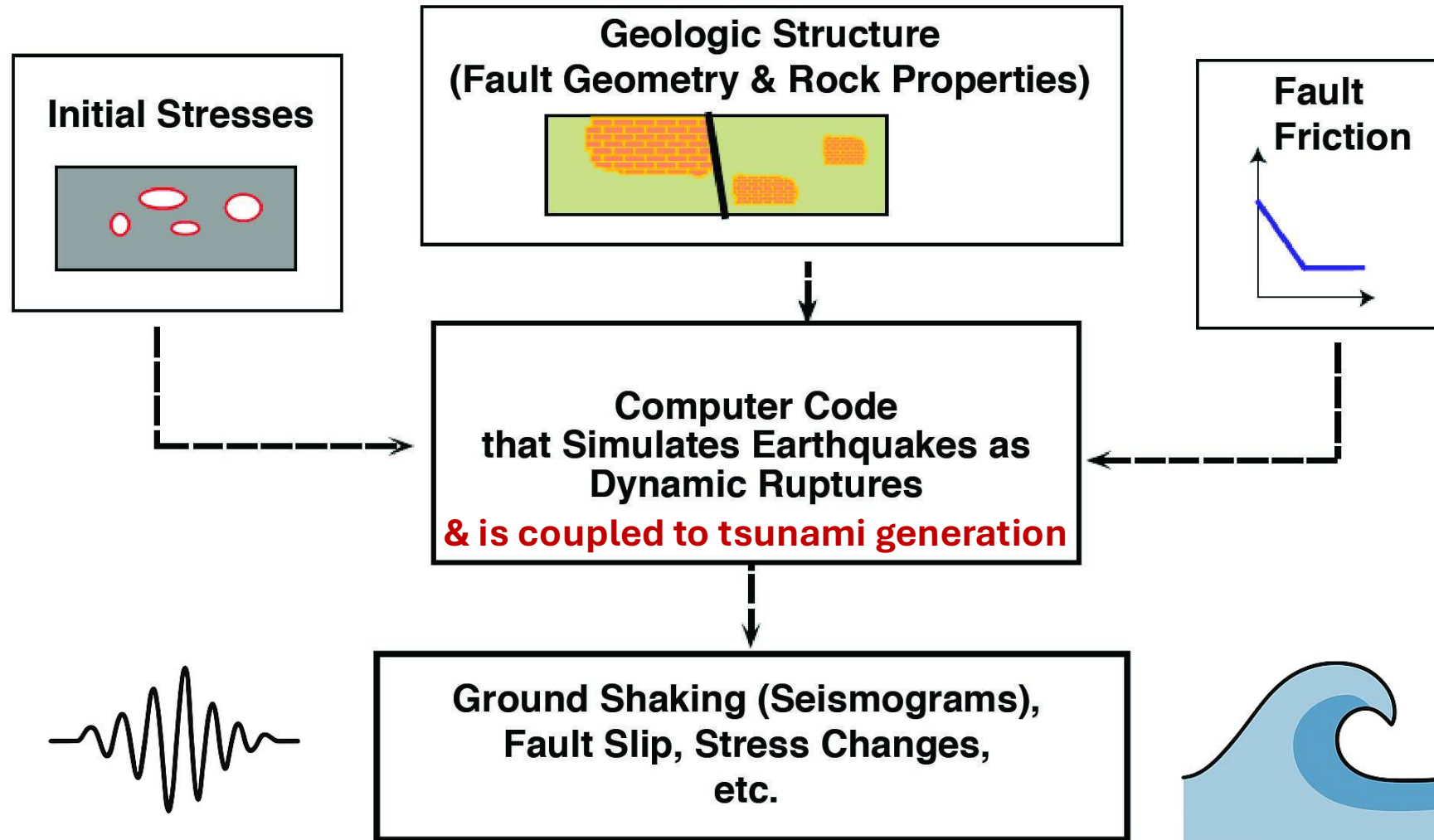


Bayesian inversion-based digital twin that employs acoustic pressure data from seafloor sensors, along with 3D coupled acoustic-gravity wave equations, to **infer earthquake-induced spatiotemporal seafloor motion** in real time and forecast tsunami propagation (Henneking et al., 2025; Glehman et al., 2025)



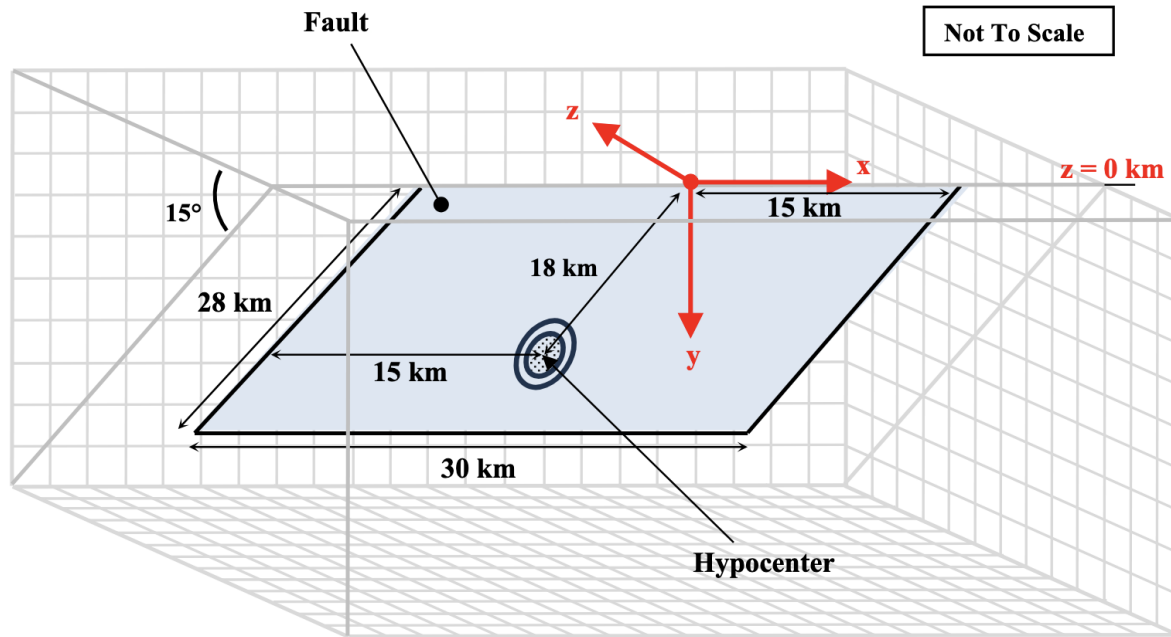
Extended **fully coupled earthquake-tsunami model** for the 2011 Mw 9.1 Tohoku-Oki earthquake with wedge inelasticity (Ma & Du, 2025)

Combined dynamic rupture and tsunami ingredients

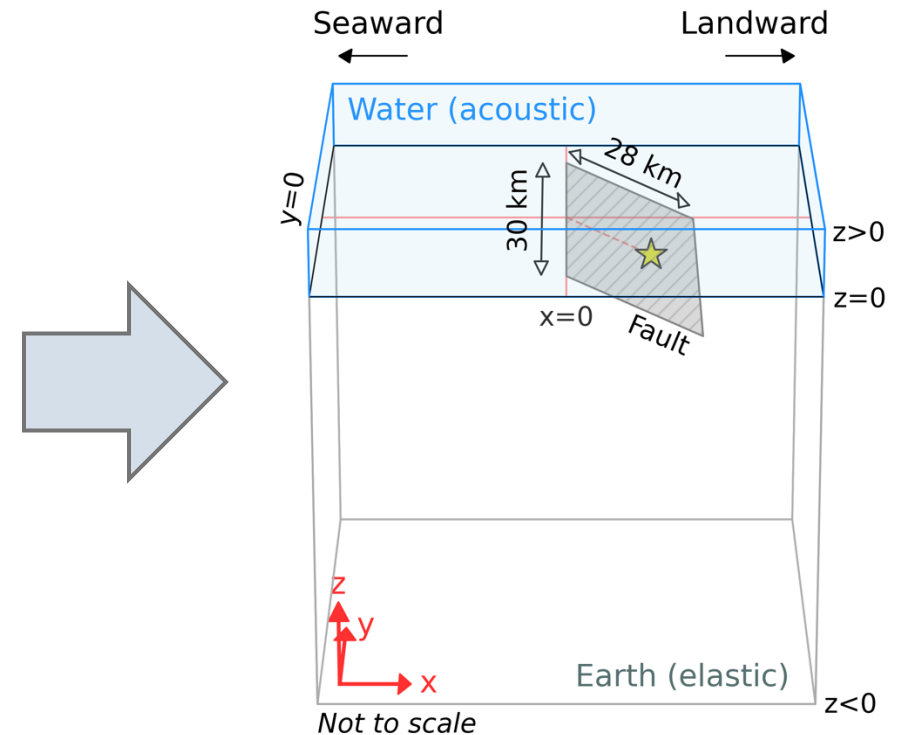


Extending an existing benchmark to tsunami generation

The Problem Version 36 (TPV36; Harris et al., in prep.)



The Tsunami Problem Version 1 (TTPV1)



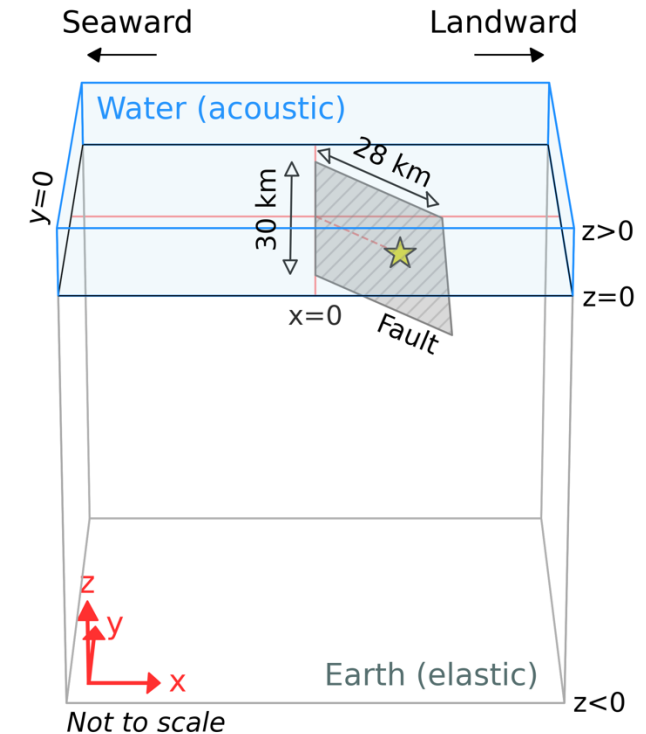
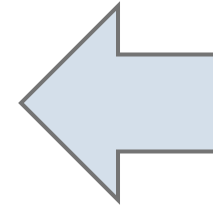
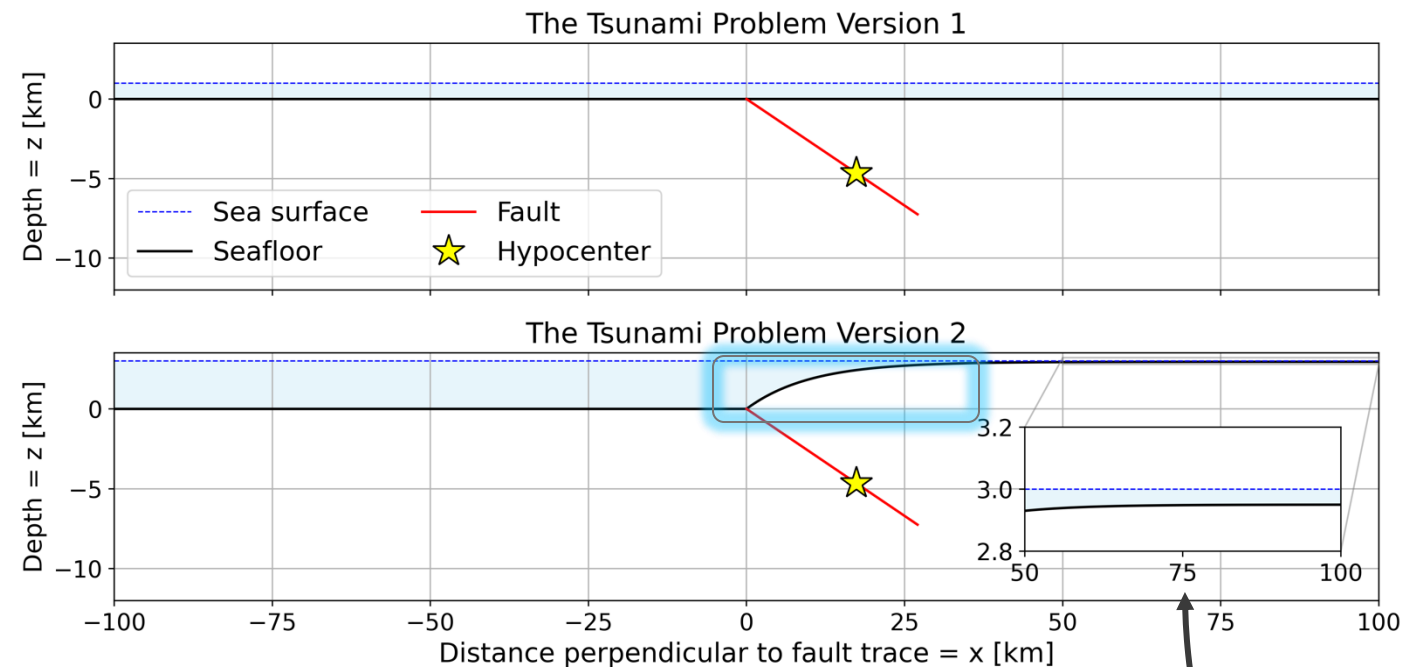
- Spontaneous 3D earthquake **dynamic rupture** on a 15° dipping thrust fault reaching the Earth's surface
- Rectangular planar fault measures 30 km along-strike and 28 km down-dip

- The **fully coupled** method combines earthquake dynamic rupture and tsunami generation into a **single simulation**
- We can capture 3D elastic, acoustic, and tsunami waves, including dispersion effects, **simultaneously** (e.g., Maeda & Furumura, 2013)

Extending an existing benchmark to tsunami generation

The Tsunami Problem Version 1 (TTPV1)

Side view:



- In TTPV1, a **1 km thick uniform** water layer is added atop the elastic medium
- A water layer with **variable depth** is atop the elastic Earth, with 3 km water depth on the seaward side, sloping to shallow depth (~50 m) landwards

- The **fully coupled** method combines earthquake dynamic rupture and tsunami generation into a **single simulation**
- We can capture 3D elastic, acoustic, and tsunami waves, including dispersion effects, **simultaneously** (e.g., Maeda & Furumura, 2013)

Continuum problem

The **fully coupled** method combines the elastic wave equation in the **solid Earth** with the linearized equations of motion governing small perturbations of a **compressible ocean** about a rest state in hydrostatic equilibrium (Lotto and Dunham 2015; Wilson and Ma, 2021; Krenz et al., 2021)

Earth

- Velocity-stress formulation of the elastodynamic equations

$$\frac{\partial \sigma_{ij}}{\partial t} - \lambda \delta_{ij} \frac{\partial v_k}{\partial x_k} - \mu \left(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i} \right) = 0$$
$$\rho \frac{\partial v_i}{\partial t} - \frac{\partial \sigma_{ij}}{\partial x_j} = 0.$$

- six stress components of stress:

$$\sigma_{xx}, \sigma_{xy}, \sigma_{xz}, \sigma_{yy}, \sigma_{yz}, \sigma_{zz}$$

- three velocity components:

$$v_x, v_y, v_z$$

- solid Earth density ρ and Lamé parameters λ, μ

Ocean

- Combine conservation of mass with a linearized equation of state
- Conservation of momentum
- Linearized free surface boundary condition at the sea surface

$$p'(x, y, H, t) = \rho_w g \eta.$$

- Linearized kinematic condition for the sea surface

$$\frac{\partial \eta}{\partial t} = v_z(x, y, H, t).$$

p' : pressure perturbation

ρ_w : water density

$H = z - \eta(x, y, t)$: height of ocean at rest

η : vertical displacement of the sea surface

Interface conditions

Which conditions have to hold at the **material interfaces**, which can generally be elastic-elastic, acoustic-acoustic, elastic-acoustic, or acoustic-elastic?

→ Unit normal n_i points from lower side (−) of interface to the upper side (+)

Elastic-Vacuum

- “Stress-free” surface

$$\sigma_{ij}^- n_i = 0$$

→ Rayleigh waves

Acoustic-Acoustic

- Continuity of normal velocity and of pressure p

$$v_i^+ n_i = v_i^- n_i$$

$$p^+ = p^-$$

Elastic-Elastic

- Continuity of velocity and traction

$$v_i^+ = v_i^-$$

$$\sigma_{ij}^+ n_i = \sigma_{ij}^- n_i$$

→ Stoneley waves

Elastic-Acoustic & Acoustic-Elastic

- Continuity of normal velocity and traction components of stress

(Sochacki et al., 1991;
Wilcox et al., 2010)

$$v_i^+ n_i = v_i^- n_i$$

$$-p' n_j = \sigma_{ij}^+ n_i = \sigma_{ij}^- n_i$$

→ Scholte waves

Code overview:

How to deal with the “trench” at the elastic-acoustic interface?

MAFE: velocity-stress-based finite element method with linear tetrahedral elements
(Ma and Nie, 2019)

drdg3d: nodal mixed-flux-based discontinuous Galerkin finite element method
(Zhang et al., 2023)

SeisSol: modal arbitrary high-order derivatives discontinuous Galerkin finite element method
(Gabriel et al., 2025)

Code overview:

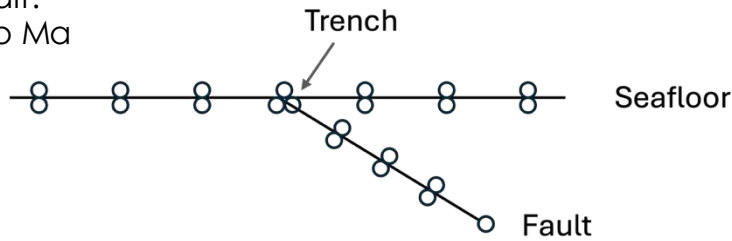
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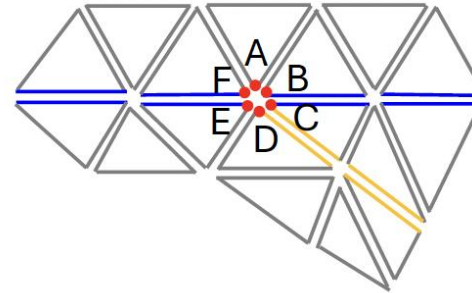
SeisSol: modal arbitrary high-order derivatives discontinuous Galerkin finite element method (Gabriel et al., 2025)

Credit:
Shuo Ma



- Split-node scheme
- Introduction of fictitious node (“center of mass”)

Credit:
Wenqiang Zhang



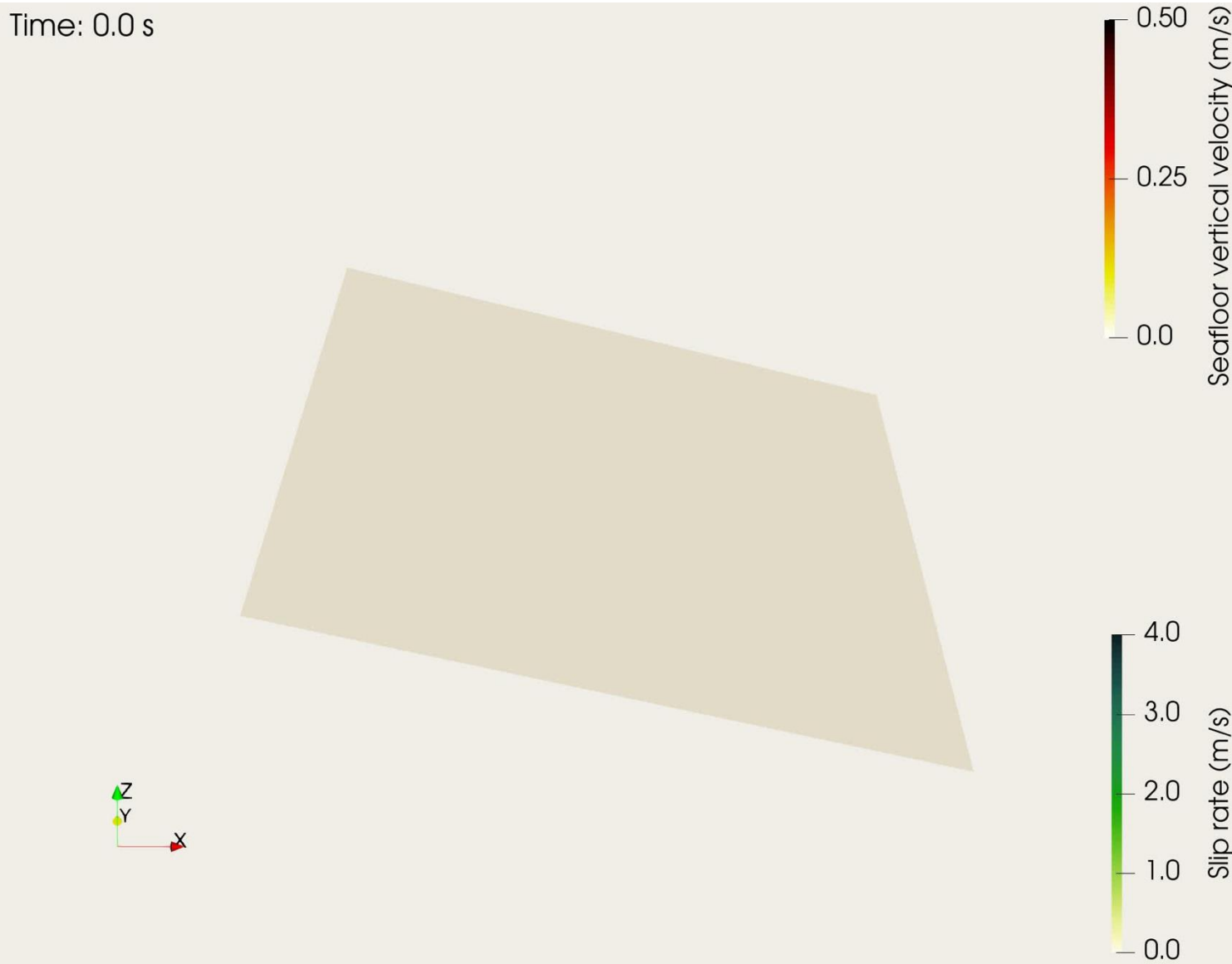
→ All boundary conditions need to be explicitly implemented (via numerical fluxes)

→ SeisSol:

- **Godunov:** solves a Riemann problem at each cell interface exactly
- **Rusanov:** approximate Riemann solver, with the averaging of fluxes from both sides of an interface

- Acoustic-acoustic: AB,AF
- Acoustic-elastic: BC,EF
- Elastic-elastic: CD,DE
- Fault: CD

Spontaneous earthquake dynamic rupture



- Simulation of a M_W 7.1 earthquake with a peak fault slip of 3.4 m to test modeling capabilities
- Linear slip-weakening friction law with (on-fault) cohesion
- Initial normal stress and the initial shear stress are proportional to distance down-dip
- Simulation time of 4 minutes

Movie: First 30 s of spontaneous dynamic rupture for SeisSol with the seafloor vertical velocity for TTPV1

TTPV1: Comparison of the final seafloor displacement

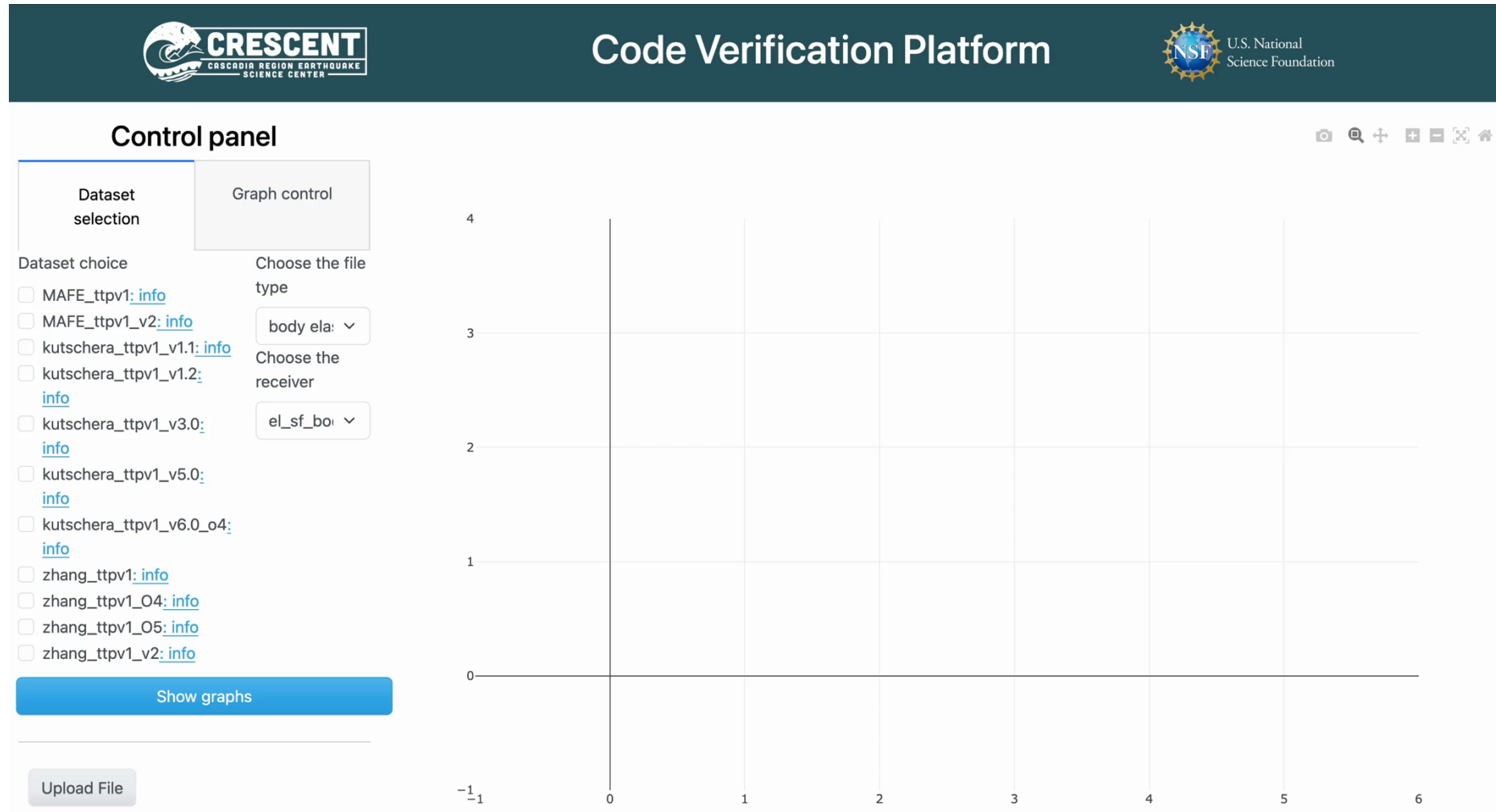
TTPV1 & 2 BENCHMARK DESCRIPTION

UPLOAD YOUR RESULTS

TTPV1 RESULTS

TTPV2 RESULTS

- Visualization based on the new code verification platform (<https://cascadiaquakes.org/det>)



Loïc Bachelot



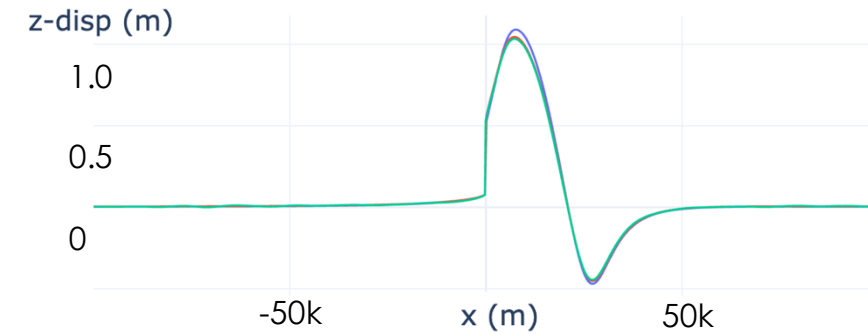
Amanda Thomas



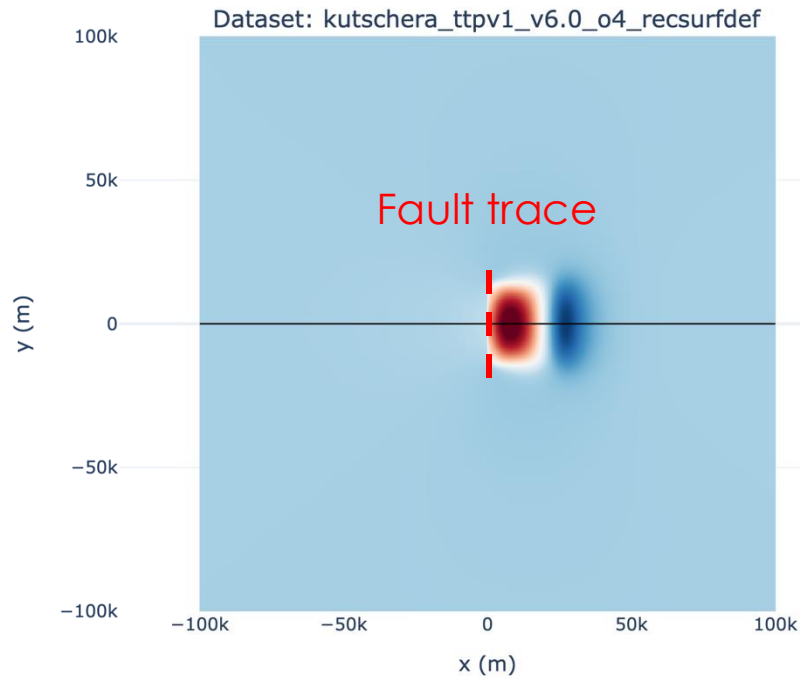
William Marfo

TTPV1: Comparison of the final seafloor displacement

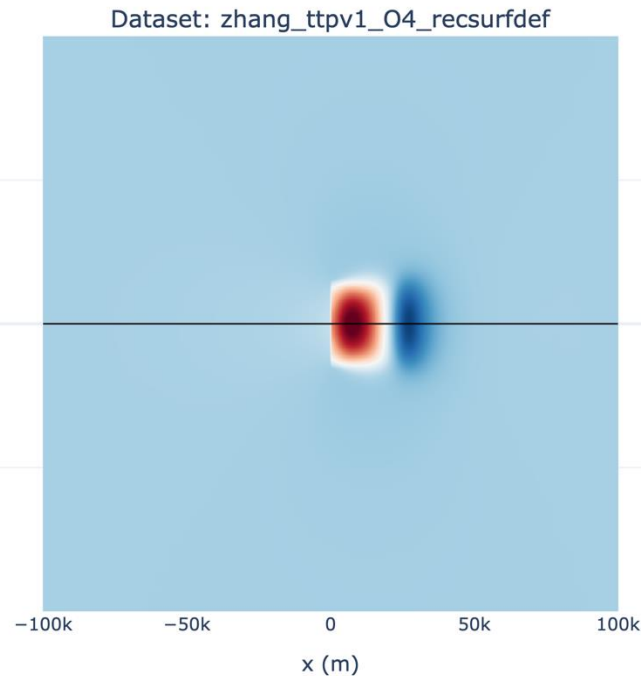
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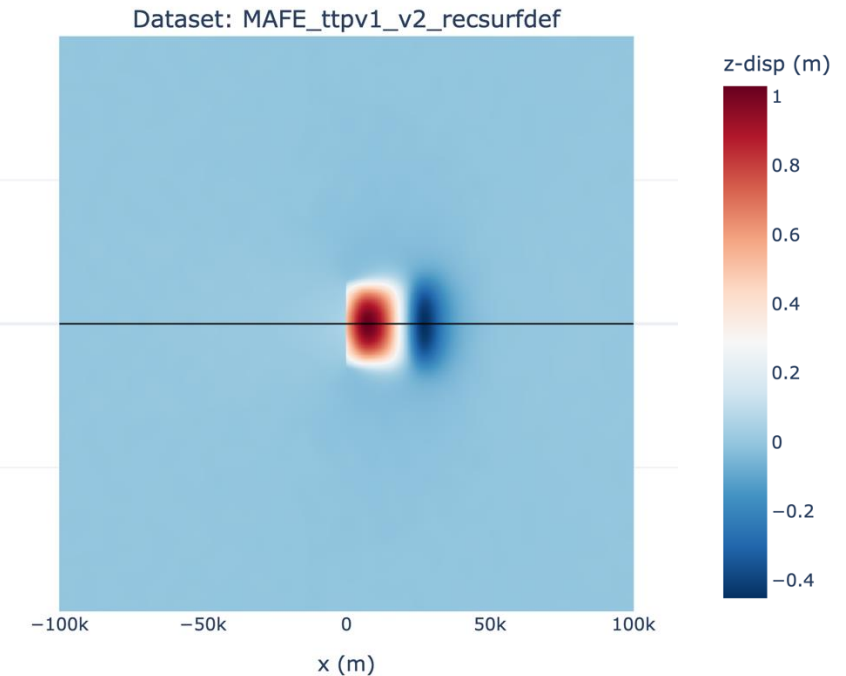
SeisSol



drdg3d



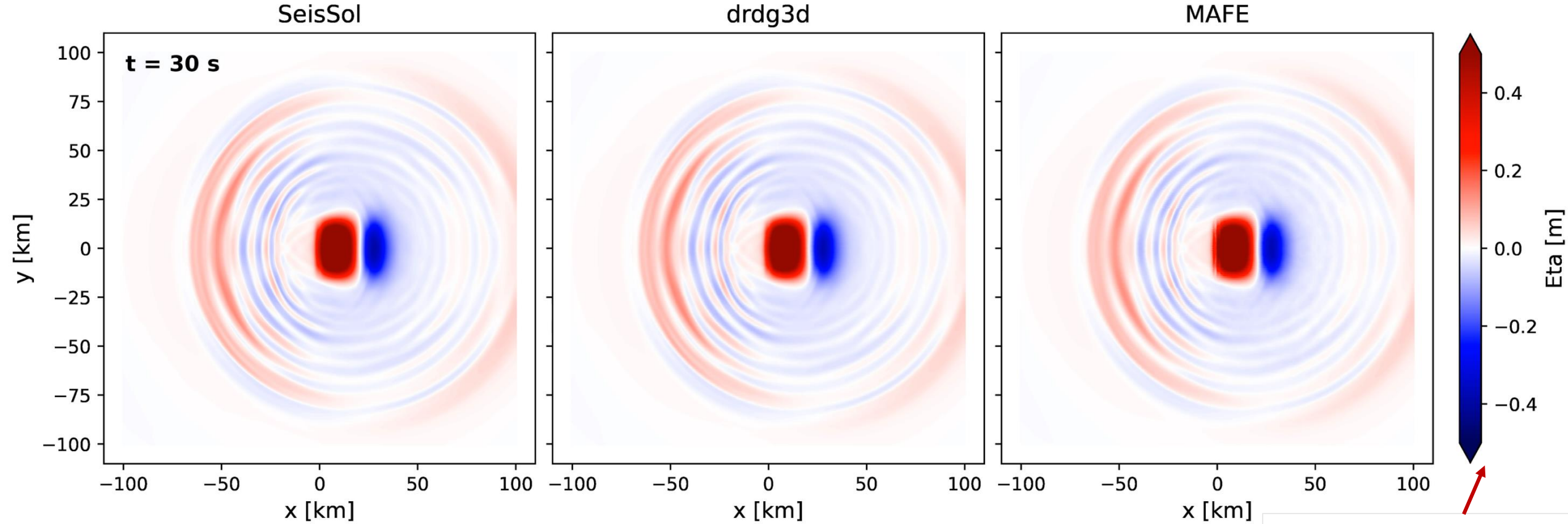
MAFE



- Distribution of spatial seafloor displacement in good agreement, with peak uplift of 1.05 ± 0.02 m, and peak subsidence of -0.45 ± 0.01 m

TTPV1: Comparison of the time-dependent tsunami generation on the sea surface

- All simulations include transient seismo-acoustic waves (visible at 30 s)

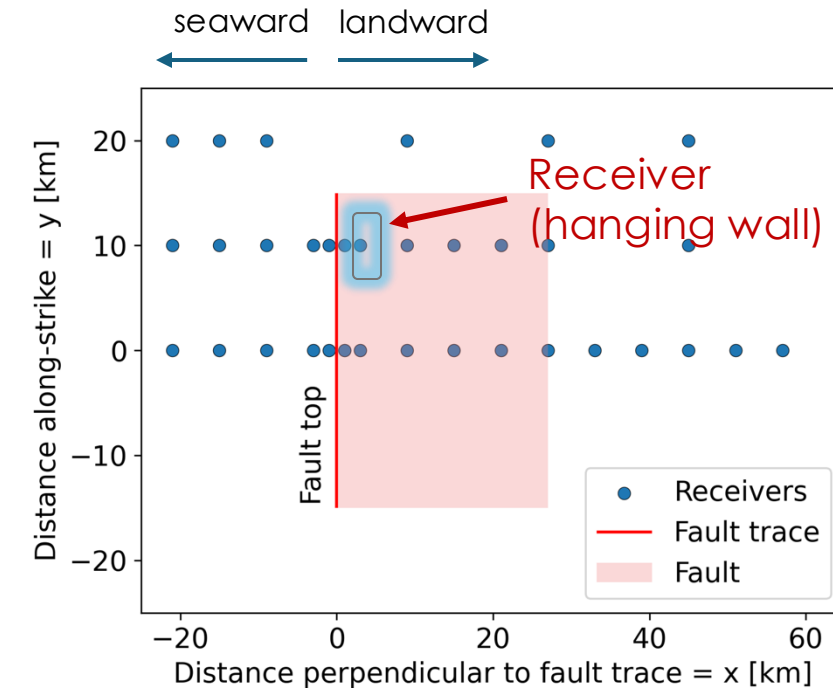
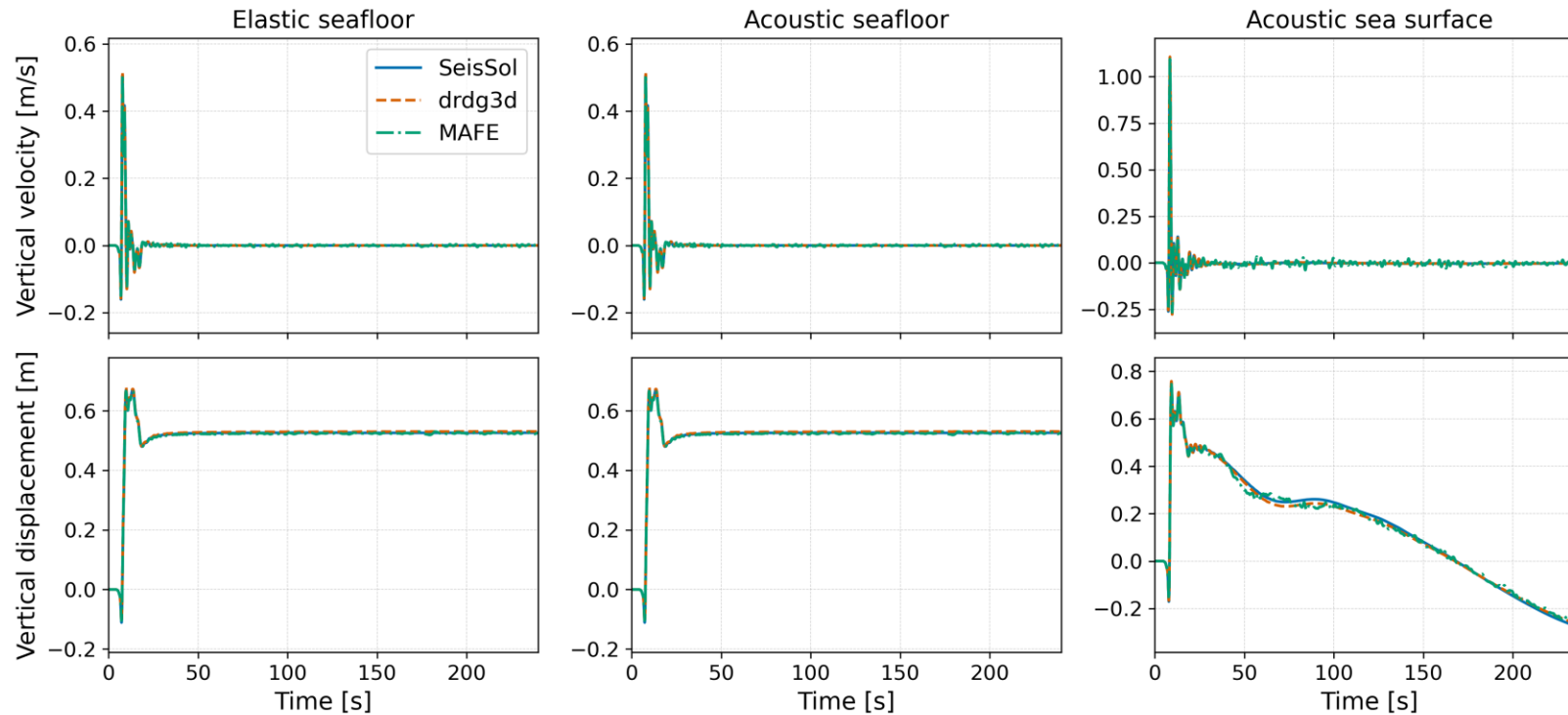


- Key features of the tsunami generation are in good agreement

Eta = vertical displacement of the **sea surface** (tsunami + seismo-acoustic waves)¹³

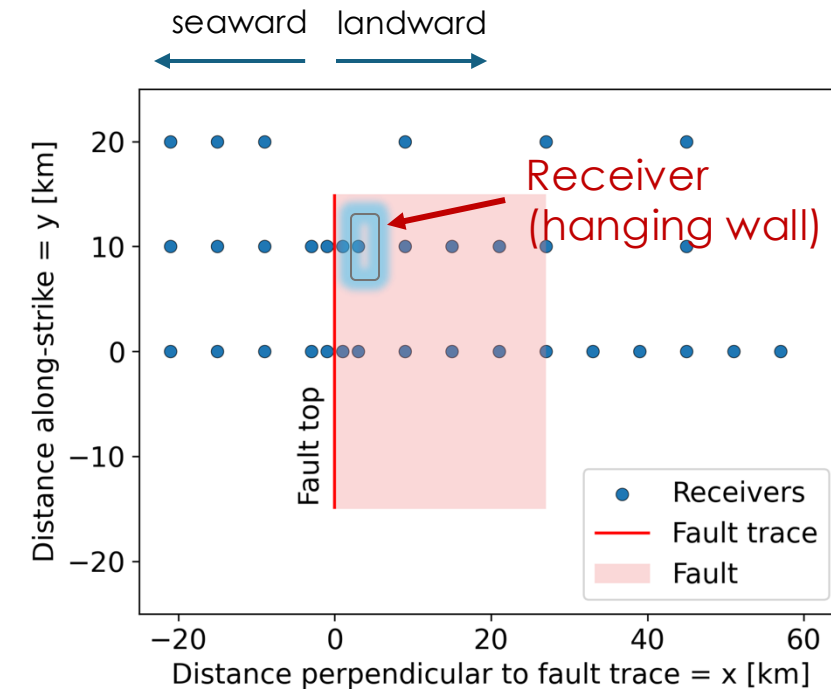
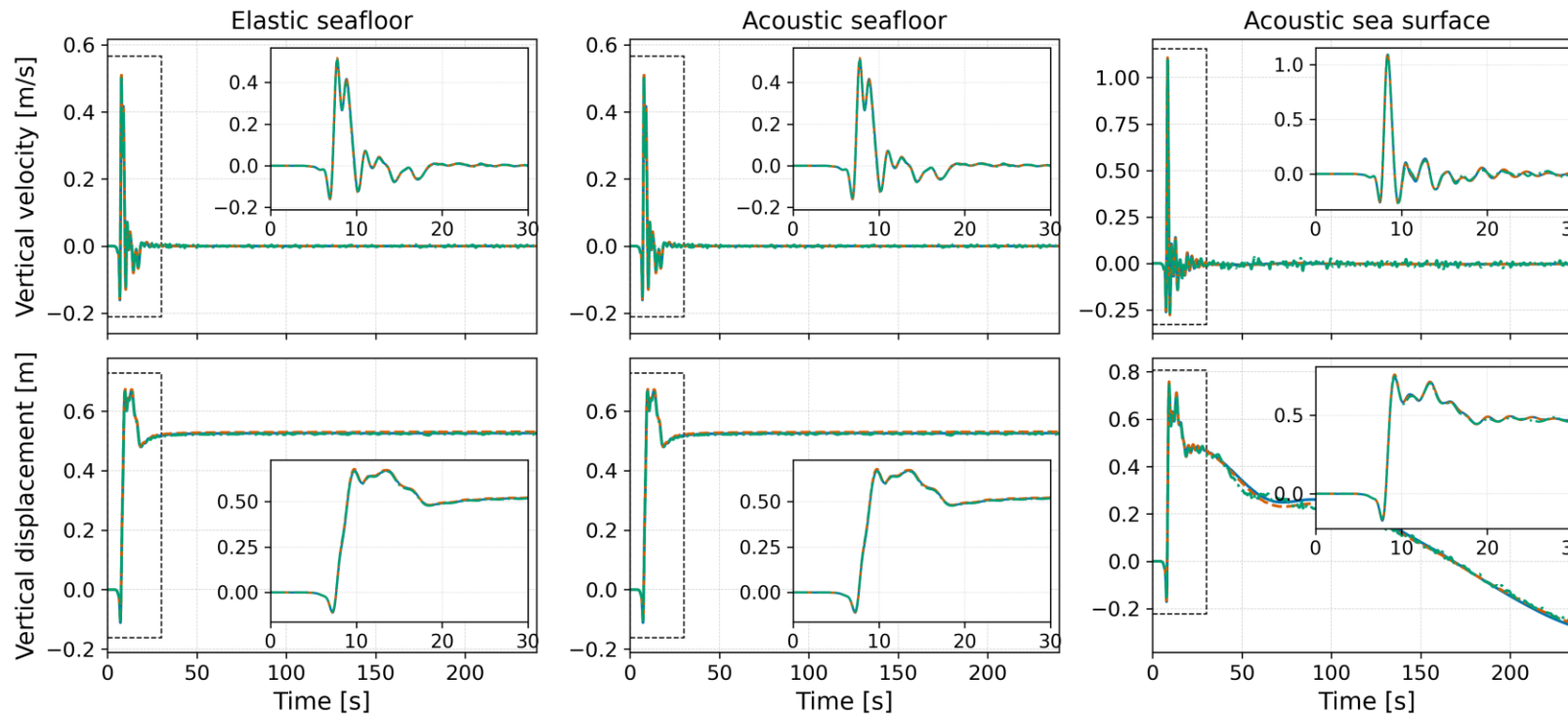
TTPV1: Comparison of the receiver time-series

- Three-component velocities and displacements over time (elastic & acoustic)



TTPV1: Comparison of the receiver time-series

- Three-component velocities and displacements over time (elastic & acoustic)



- Great agreement of the initial generation phase (first 30 s), clean of any numerical noise

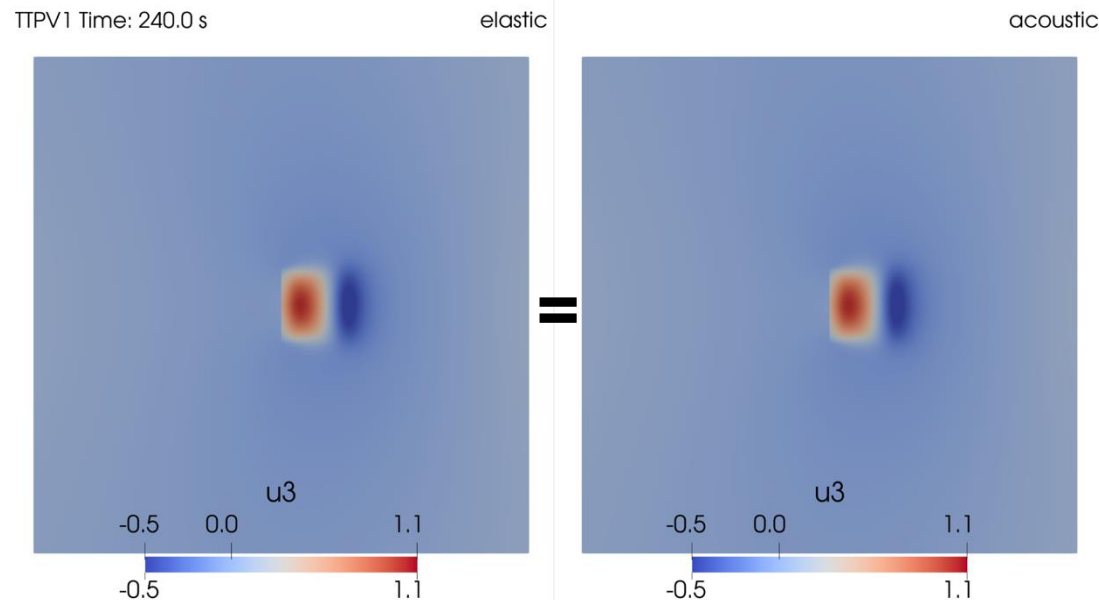
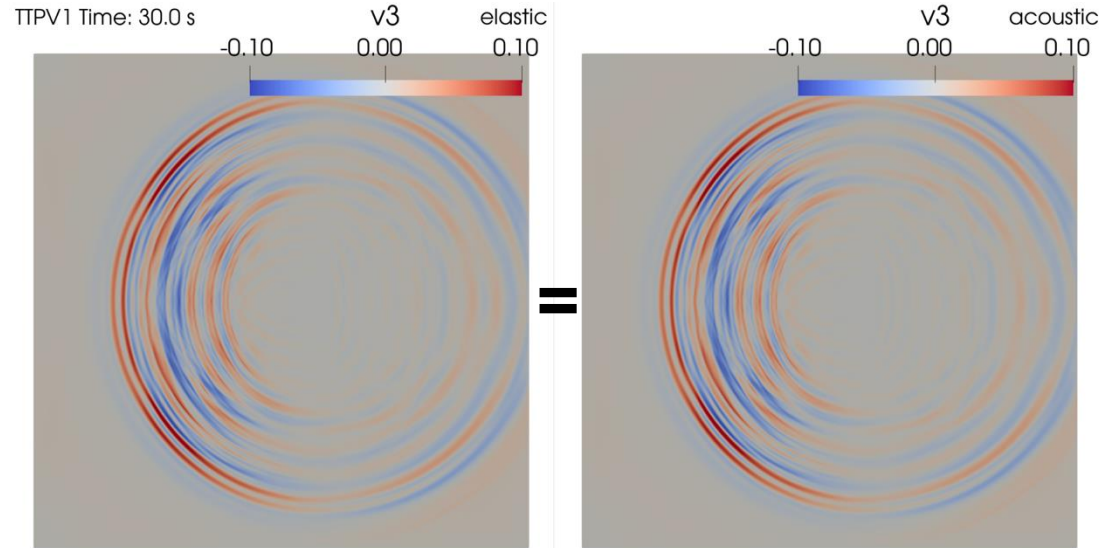
Seafloor continuity requirement in TTPV1

Recall continuity of normal velocity at elastic-acoustic interface :

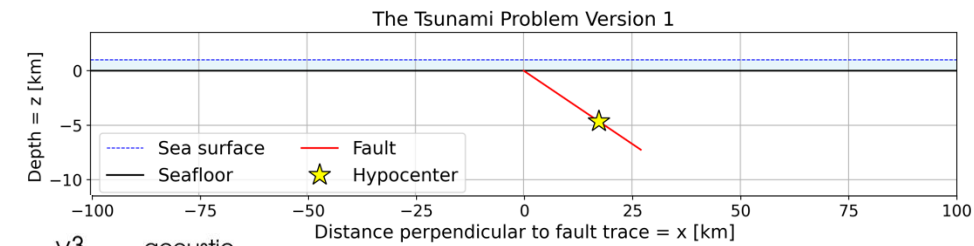
$$v_i^+ n_i = v_i^- n_i$$

Since TTPV1 has a flat bathymetry:

$$v_z = v_i^+ n_i$$

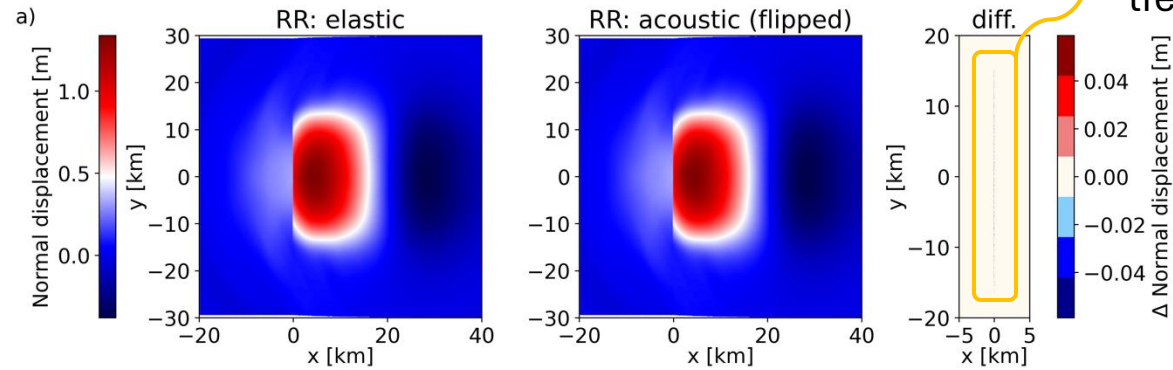


SeisSol

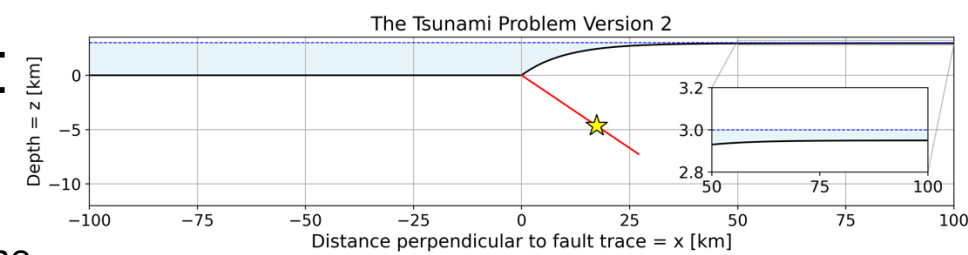


Accuracy of interface condition enforcement is affected by choice of numerical fluxes

SeisSol

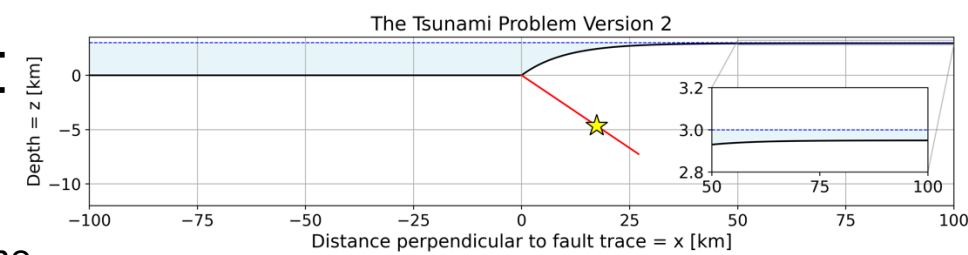


Can we do better and reduce the “trench artifacts”?



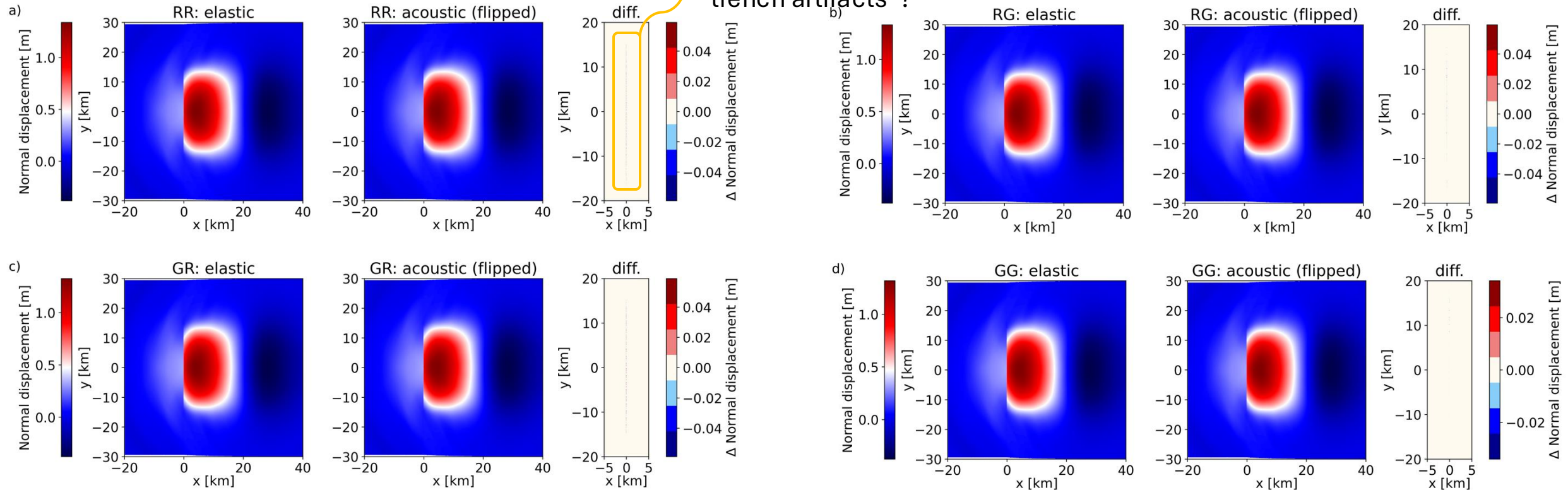
- We can use different numerical fluxes, **Rusanov (R)** and **Godunov (G)**, for solving wave propagation (i) in the volume and (ii) for faces of cells adjacent to the fault

Accuracy of interface condition enforcement is affected by choice of numerical fluxes



SeisSol

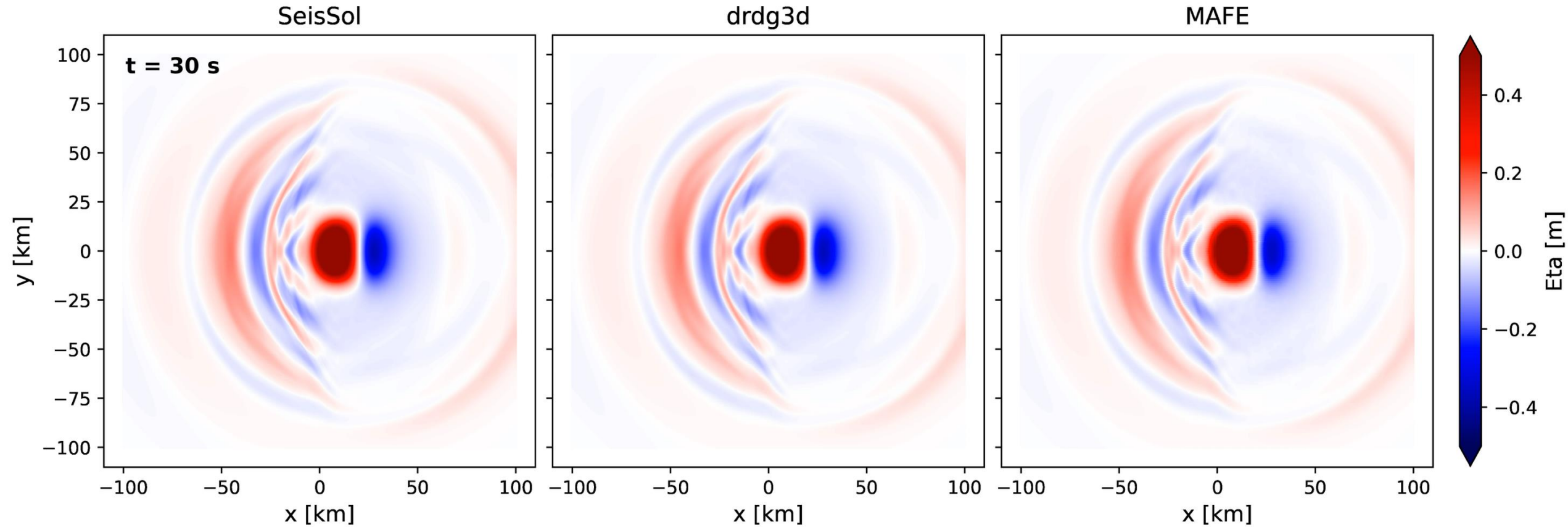
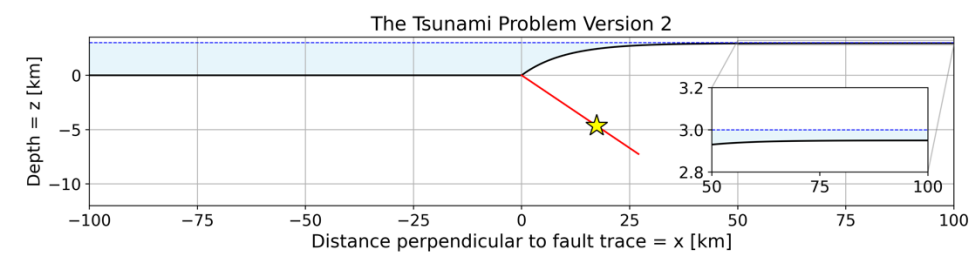
Can we do better and reduce the “trench artifacts”?



- We can use different numerical fluxes, **Rusanov (R)** and **Godunov (G)**, for solving wave propagation (i) in the volume and (ii) for faces of cells adjacent to the fault
- When using GG, the peak absolute amplitude of trench artifacts is reduced by 42.2%, 42.8%, and 42.2% when compared against GR, RG, and RR fluxes, respectively

TTPV2: Comparison of the time-dependent tsunami generation on the sea surface

- All simulations include transient seismo-acoustic waves (visible at 30 s)



- Key features of the tsunami generation are in good agreement

Conclusions and outlook

- **3 codes were verified in 2 benchmarks** to model 3D fully coupled earthquake dynamic rupture and tsunami generation
- Good agreement of key features in the initial tsunami generation for both TTPV1 & TTPV2
- Numerical fluxes affect the accuracy of enforcing elastic–acoustic interface conditions at complex geometries
- Size matters → sufficiently large domain size is required for the benchmarks
- What's next for the SCEC-USGS Dynamic Rupture Verification Group?
 - Explore other, more complex/realistic topography?
 - Should we include off-fault plasticity again?

Benchmark
description:

<https://doi.org/10.5281/zenodo.15389414>

